

Correcting 4 False Assumptions
in PG&E's Seismic Source Characterization [2015]
and Update [2024] that Caused PG&E to
**Seriously Underestimate Seismic Hazard
at Diablo Canyon Nuclear Power Plant**

2024.07.17 presentation to the Petition Review Board of the
U.S. Nuclear Regulatory Commission

by Peter Bird, Professor Emeritus, UCLA
consulting to San Luis Obispo Mothers for Peace,
Friends of the Earth, & Environmental Working Group



4 False Assumptions

- #1. The Irish Hills are uplifting as a rigid block, with no internal deformation.
- #2. Active thrust faults may dip at any angle.
- #3. Geologic structures older than ~ 0.33 Ma are irrelevant to seismic hazard estimation.
- #4. GPS geodetic velocities are not useful for site-specific seismic hazard estimation.

#1. The Irish Hills are uplifting (at ~0.2 mm/year, stipulated) as a rigid block, with no internal deformation.

Therefore, thrust faulting occurs only at the margins (Los Osos thrust, San Luis Bay thrust, **?Inferred Coastline thrust?**) with fault throw (vertical) rates of ~0.2 mm/year.

HOWEVER:

- The geologic map shows tight folding of Late Miocene sedimentary rocks has occurred since 6~5 Ma. Therefore, the Irish Hills are not rigid, and additional blind thrust faults are active in the interior.
- Rigid-body uplift does not produce crustal thickening. Therefore, if the Irish Hills were a rigid block, they would have a positive isostatic gravity anomaly. However, data shows a negative isostatic gravity anomaly, indicating more than simple Airy compensation by crustal roots (more than the typical Airy ratio of 6:1).

THEREFORE:

A simple isostatic model for the total rate of thrust-fault slip under the Irish Hills is at least:

$$(0.2 \text{ mm/year uplift}) \times 6 / \sin(25^\circ \text{ dips}) = 2.8 \text{ mm/year}$$

This is the 1st of 3 independent analytic estimates developed in this presentation.



Basemap excerpted from Plate 7-1 of Pacific Gas & Electric (2015) Seismic Source Characterization for Diablo Canyon Power Plant; Inferred Coastline thrust fault trace (in red) added by P. Bird, 2024.06.19. Triangle symbols show direction of dip of thrust faults.

#1. The Irish Hills are uplifting (at ~0.2 mm/year, stipulated) as a rigid block, with no internal deformation.

Therefore, thrust faulting occurs only at the margins (Los Osos thrust, San Luis Bay thrust, **?Inferred Coastline thrust?**) with fault throw (vertical) rates of ~0.2 mm/year.



HOWEVER:

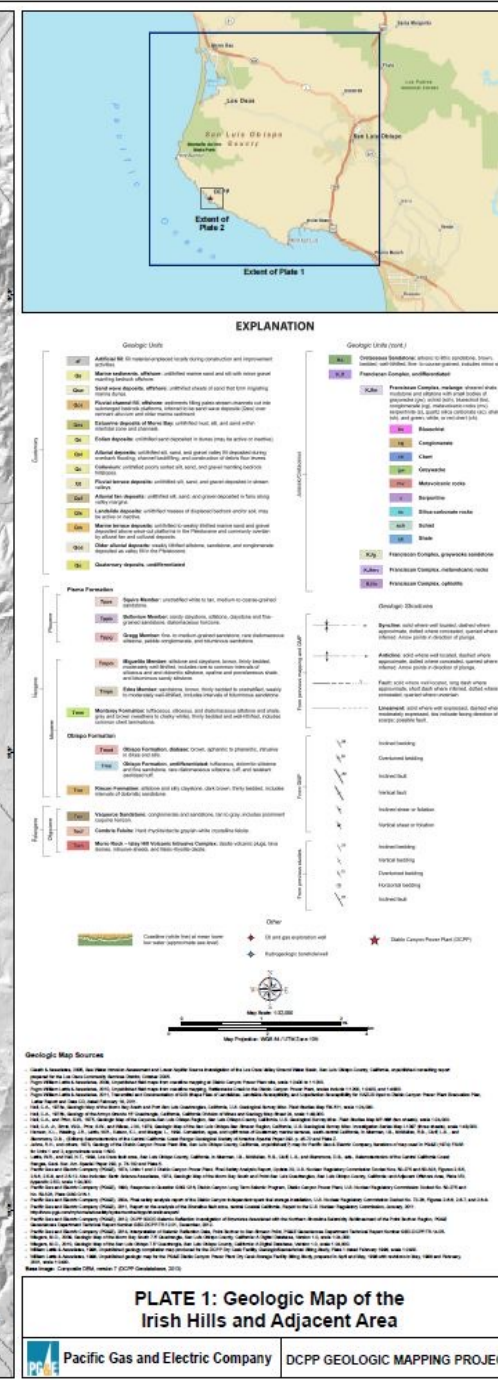
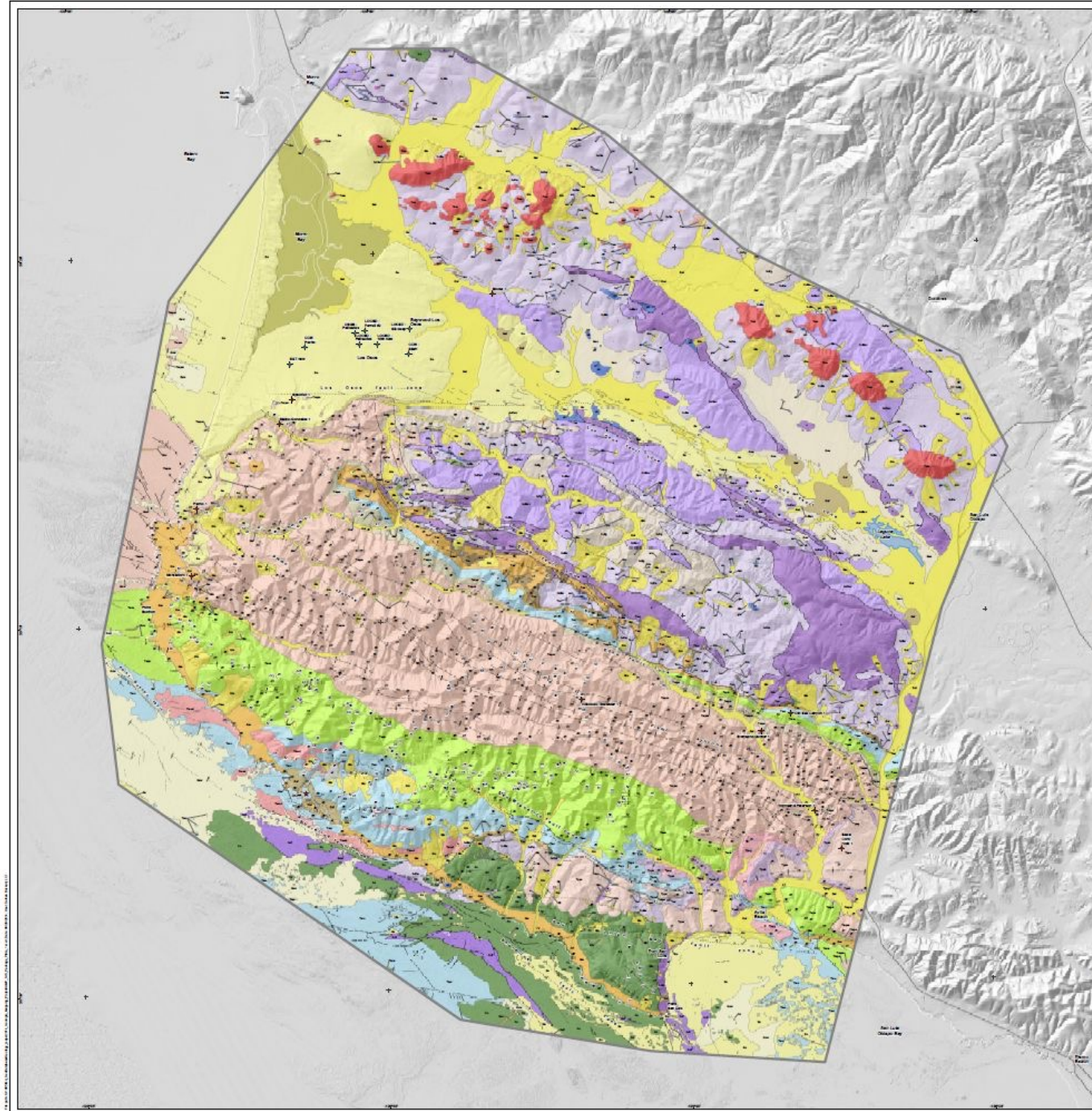
- The geologic map shows tight folding of Late Miocene sedimentary rocks has occurred since 5 Ma. Therefore, the Irish Hills are not rigid, and additional blind thrust faults are active in the interior.
- Rigid-body uplift does not produce crustal thickening. Therefore, if the Irish Hills were a rigid block, they would have a positive isostatic gravity anomaly. However, data shows a negative isostatic gravity anomaly, indicating more than simple Airy compensation by crustal roots (more than the typical Airy ratio of 6:1).

THEREFORE:

A simple isostatic model for the total rate of thrust-fault slip under the Irish Hills is at least:

$$(0.2 \text{ mm/year uplift}) \times 6 / \sin(25^\circ \text{ dips}) = 2.8 \text{ mm/year}$$

This is the 1st of 3 independent analytic estimates developed in this presentation.



The geologic map of the Irish Hills demonstrates large internal deformation since 5 Ma, especially in the Pismo Syncline.

#1. The Irish Hills are uplifting (at ~0.2 mm/year, stipulated) as a rigid block, with no internal deformation.

Therefore, thrust faulting occurs only at the margins (Los Osos thrust, San Luis Bay thrust, ?Inferred Coastline thrust?) with fault throw (vertical) rates of ~0.2 mm/year.

HOWEVER:

➤ The geologic map shows tight folding of Late Miocene sedimentary rocks has occurred since 5 Ma. Therefore, the Irish Hills are not rigid, and additional blind thrust faults are active in the interior.



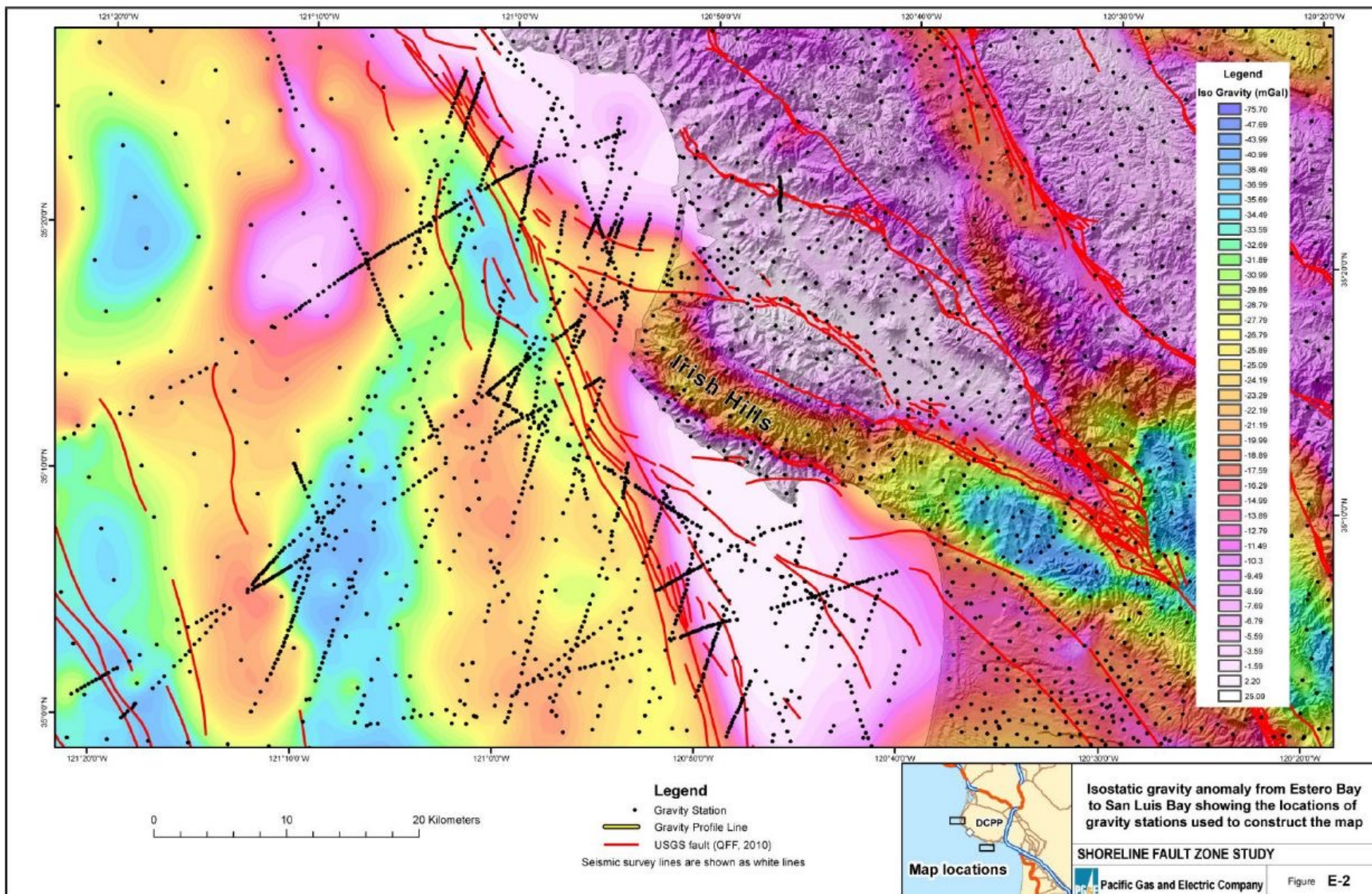
➤ Rigid-body uplift does not produce crustal thickening. Therefore, if the Irish Hills were a rigid block, they would have a positive isostatic gravity anomaly. However, data shows a negative isostatic gravity anomaly, indicating more than simple Airy compensation by crustal roots (more than the typical Airy ratio of 6:1).

THEREFORE:

A simple isostatic model for the total rate of thrust-fault slip under the Irish Hills is at least:

$$(0.2 \text{ mm/year uplift}) \times 6 / \sin(25^\circ \text{ dips}) = 2.8 \text{ mm/year}$$

This is the 1st of 3 independent analytic estimates developed in this presentation.



The negative isostatic gravity anomaly here means that:

The topography of the Irish Hills is not just isostatically compensated, it is OVER-compensated by crustal thickening.

Figure 6-2. Large-Scale Residual Isostatic Gravity Anomaly Map Showing a Negative Gravity Anomaly Coincident with the Irish Hills (modified from Langenheim et al., 2008 and PG&E, 2011, Figure E-2)

#1. The Irish Hills are uplifting (at ~0.2 mm/year, stipulated) as a rigid block, with no internal deformation.

Therefore, thrust faulting occurs only at the margins (Los Osos thrust, San Luis Bay thrust, **?Inferred Coastline thrust?**) with fault throw (vertical) rates of ~0.2 mm/year.

HOWEVER:

- The geologic map shows tight folding of Late Miocene sedimentary rocks has occurred since 6~5 Ma. Therefore, the Irish Hills are not rigid, and additional blind thrust faults are active in the interior.
- Rigid-body uplift does not produce crustal thickening. Therefore, if the Irish Hills were a rigid block, they would have a positive isostatic gravity anomaly. However, data shows a negative isostatic gravity anomaly, indicating more than simple Airy compensation by crustal roots (more than the typical Airy ratio of 6:1).

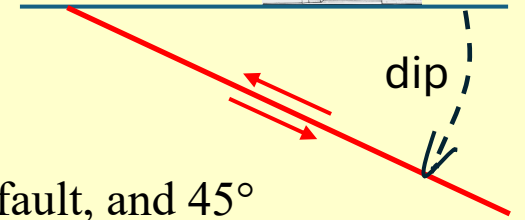
THEREFORE:



A simple isostatic model for the total rate of thrust-fault slip under the Irish Hills is at least:

$$(0.2 \text{ mm/year uplift}) \times 6 / \sin(25^\circ \text{ dips}) = 2.8 \text{ mm/year}$$

This is the 1st of 3 independent analytic estimates developed in this presentation.



#2. Active thrust faults may dip at any angle (measured from the horizontal).

PG&E assigned alternative model dips of 30°, 50°, and 80° for the Los Osos thrust fault, and 45° to 75° for the San Luis Bay thrust fault.

HOWEVER:

125-year-old Mohr/Coulomb friction theory shows that thrusts never form at dips steeper than 45°, and most commonly dip at ~25° [for rock friction coefficient of 0.85; *Byerlee*, 1978].

THEREFORE:

Seismic potency rate (per m of fault trace) is defined as = (slip rate) × (down-dip width).

This important measure of earthquake generation varies as $1/\sin^2(\text{dip})$ when throw-rate is held constant (as in these 2 SSC studies).

Compared to reasonable estimates (obtained with dip of 25°), an assignment of 50° dip reduces seismic potency rate by a factor of **3.3×**. An assignment of 80° dip reduces seismic potency rate by factor of **5.4×**.

Thus, PG&E underestimated seismic potency of these 2 thrusts (which were the only ones they recognized) by large factors.

MAXIMUM FRICTION

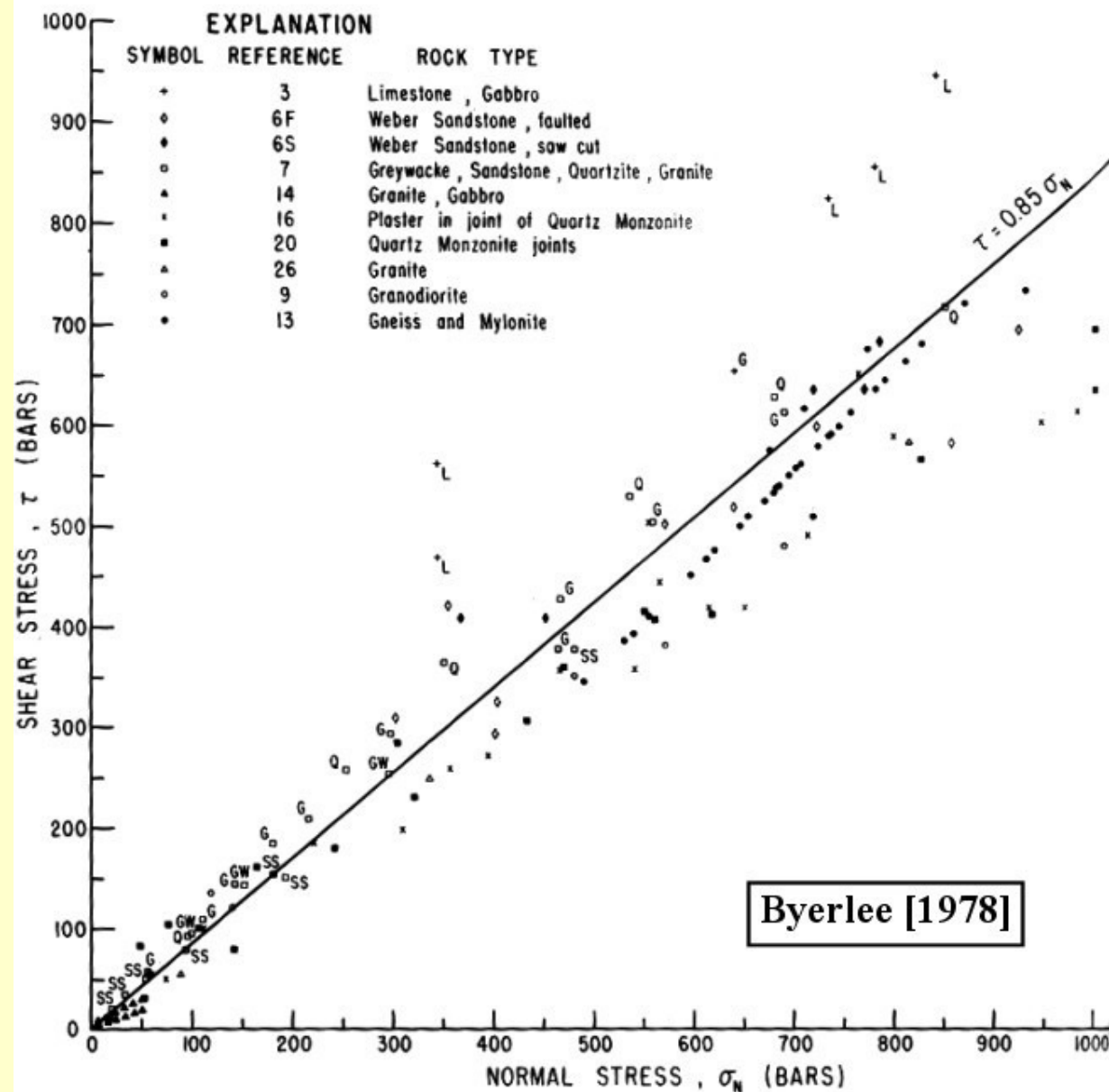
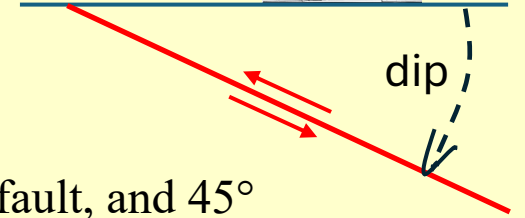


Figure 5
Shear stress plotted as a function of normal stress at the maximum friction for a variety of rock types at normal stresses to 1000 bars.



#2. Active thrust faults may dip at any angle (measured from the horizontal).

PG&E assigned alternative model dips of 30°, 50°, and 80° for the Los Osos thrust fault, and 45° to 75° for the San Luis Bay thrust fault.

HOWEVER:

125-year-old Mohr/Coulomb friction theory shows that thrusts never form at dips steeper than 45°, and most commonly dip at ~25° [for rock friction coefficient of 0.85; *Byerlee, 1978*].

THEREFORE:



Seismic potency rate (per m of fault trace) is defined as = (slip rate) × (down-dip width).

This important measure of earthquake generation varies as $1/\sin^2(\text{dip})$ when throw-rate is held constant (as in these 2 SSC studies).

Compared to reasonable estimates (obtained with dip of 25°), an assignment of 50° dip reduces seismic potency rate by a factor of **3.3×**. An assignment of 80° dip reduces seismic potency rate by factor of **5.4×**.

Thus, PG&E underestimated seismic potency of these 2 thrusts (which were the only ones they recognized) by large factors.

#3. Geologic structures older than ~0.33 Ma are irrelevant to seismic hazard estimation.

PG&E based the throw-rates of the San Luis Bay thrust fault and the Los Osos thrust fault on vertical offsets of marine & fluvial terraces with Upper Pleistocene ages, typically ~0.12 Ma.

PG&E never attempted to model the uplift and folding of sedimentary rocks in the Irish Hills which occurred since 5 Ma.

HOWEVER:

A statistical study of all dated fault offsets in California by *Bird* [2007] showed that the risk of “inapplicability to neotectonics” is constant for offset features with ages of to 3 Ma, and then rises only modestly for features of 5 Ma age [his Figure 8].

Bird [2007] also showed that a well-constrained fault offset rate requires 4~7 offset features, not just 1 or 2 [his Figure 9].

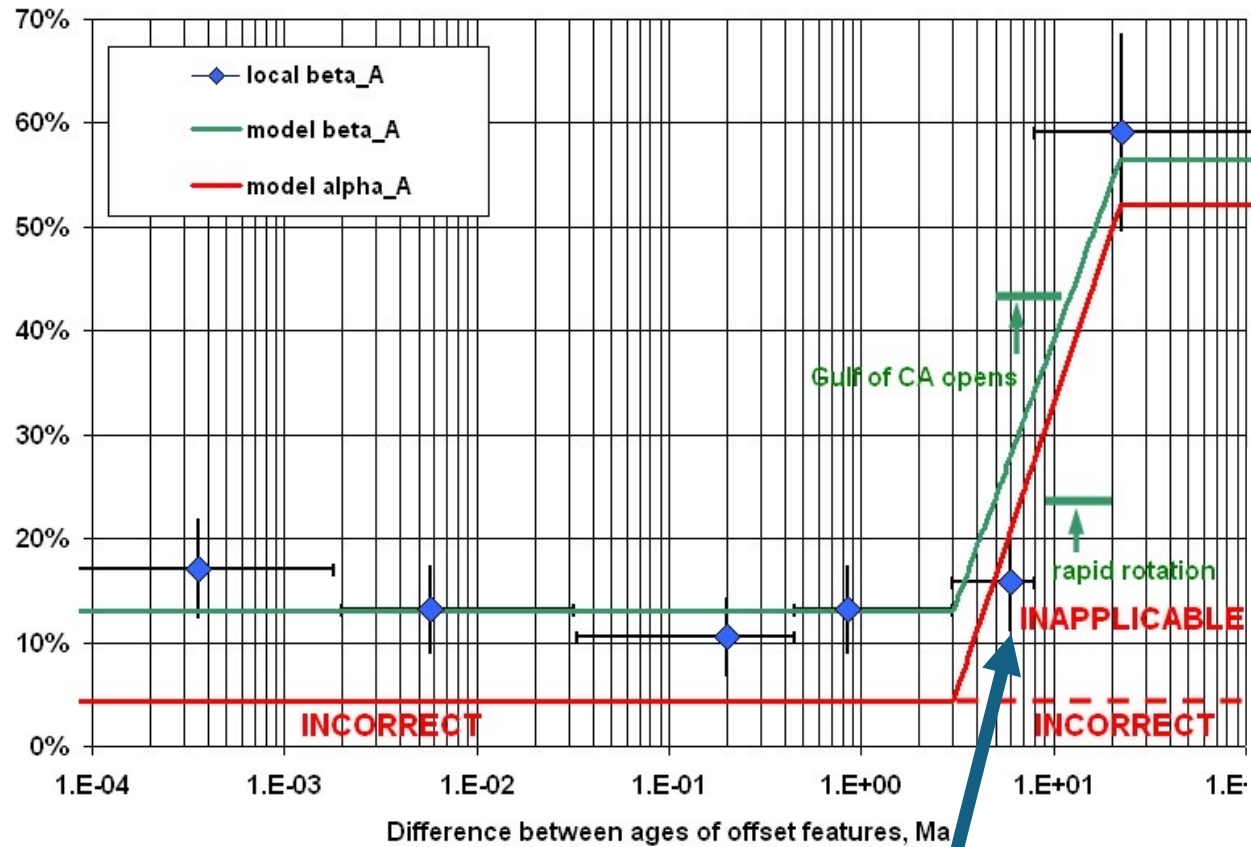
THEREFORE:

Therefore, all the structures in the Irish Hills, which formed since 5 Ma, should have been studied and modeled to provide geologic constraints on the rates of thrust-faulting.

I provided one example in Figure 1 of my March 2024 Declaration: Throw of the Obispo Formation at the San Luis Bay-Inferred Coastline thrust fault is 1.6~2.2 km since 5 Ma, implying throw-rate of 0.32~0.44 mm/year, and fault slip rate of 0.76~1.04 mm/year.

If thrusting in the Irish Hills has been symmetrical(?), then a minimum total thrust slip-rate by this method would be 1.52~2.08 mm/year. (However, this neglects any internal blind thrusts.)

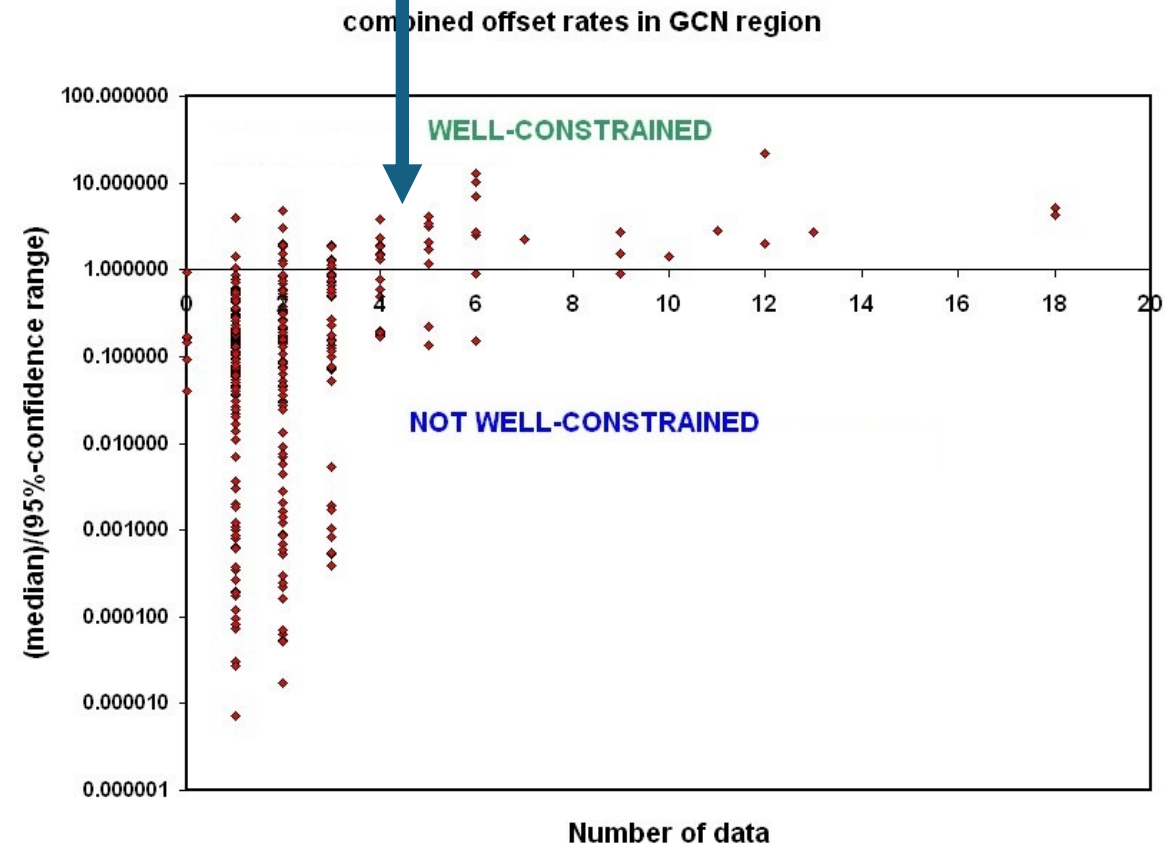
A/A comparisons in California (WGCEP data)



In California, “inapplicability to neotectonics” is only a problem for offset features older than 5 Ma. All younger features are equally relevant.

Bird [2007, Geosphere]

It takes more than 1 (or 2) offset features to give a well-constrained fault slip rate; actually it takes more than 4!



#3. Geologic structures older than ~0.33 Ma are irrelevant to seismic hazard estimation.

PG&E based the throw-rates of the San Luis Bay thrust fault and the Los Osos thrust fault on vertical offsets of marine & fluvial terraces with Upper Pleistocene ages, typically ~0.12 Ma.

PG&E never attempted to model the uplift and folding of sedimentary rocks in the Irish Hills which occurred since 5 Ma.

HOWEVER:

A statistical study of all dated fault offsets in California by *Bird* [2007] showed that the risk of “inapplicability to neotectonics” is constant for offset features with ages of to 3 Ma, and then rises only modestly for features of 5 Ma age [his Figure 8].

Bird [2007] also showed that a well-constrained fault offset rate requires 4~7 offset features, not just 1 or 2 [his Figure 9].

 THEREFORE:

Therefore, all the structures in the Irish Hills, which formed since 5 Ma, should have been studied and modeled to provide geologic constraints on the rates of thrust-faulting.

I provided one example in Figure 1 of my March 2024 Declaration: Throw of the Obispo Formation at the San Luis Bay-Inferred Coastline thrust fault is 1.6~2.2 km since 5 Ma, implying throw-rate of 0.32~0.44 mm/year, and fault slip rate of 0.76~1.04 mm/year.

If thrusting in the Irish Hills has been symmetrical(?), then a minimum total thrust slip-rate by this method would be 1.52~2.08 mm/year. (However, this neglects any internal blind thrusts.)

Mark-up (in red) by P. Bird, 2023 of

Figure 13-17 of SSC for DCPD
(PG&E, 2015)

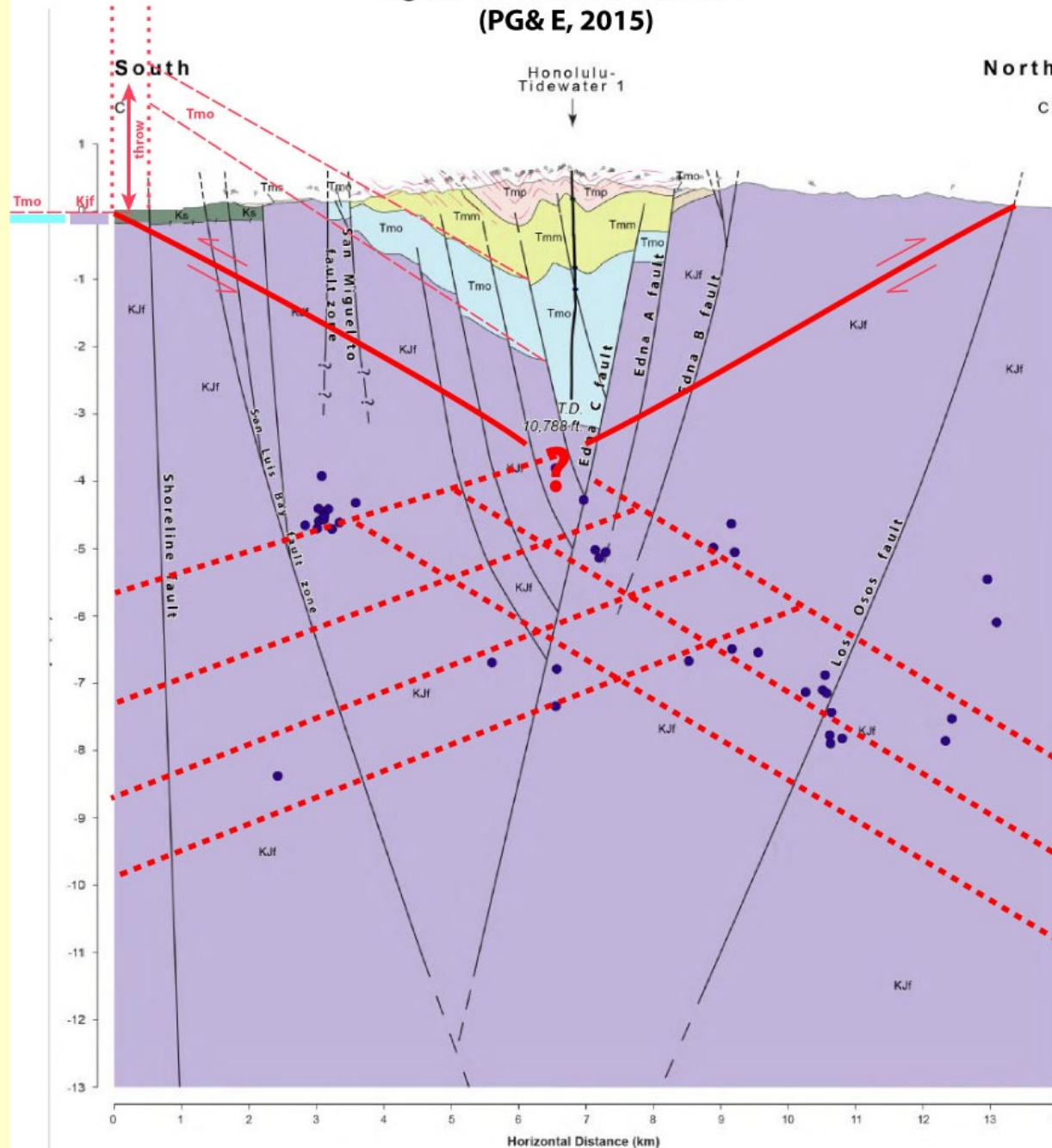


Figure 1. Revised geologic section through the Irish Hills near DCPD. The base for this figure is Figure 13-17 of the Seismic Source Characterization for DCPD (PG&E, 2015). Note that the fault dips suggested by black lines in their figure were not based on data, but were constrained by PG&E's (2015) *a priori* assumption that only strike-slip tectonics is active in the area. In red, I have suggested more plausible 25° dips for the Los Osos thrust (at right/North) and the Inferred Coastline thrust (at left/South). The upper- left portion of this figure is also edited to show the throw (vertical offset) of map unit Tmo across the Inferred Coastline thrust, discussed in my text paragraph IV.B.25(b).

#3. Geologic structures older than ~0.33 Ma are irrelevant to seismic hazard estimation.

PG&E based the throw-rates of the San Luis Bay thrust fault and the Los Osos thrust fault on vertical offsets of marine & fluvial terraces with Upper Pleistocene ages, typically ~0.12 Ma.

PG&E never attempted to model the uplift and folding of sedimentary rocks in the Irish Hills which occurred since 5 Ma.

HOWEVER:

A statistical study of all dated fault offsets in California by *Bird* [2007] showed that the risk of “inapplicability to neotectonics” is constant for offset features with ages of to 3 Ma, and then rises only modestly for features of 5 Ma age [his Figure 8].

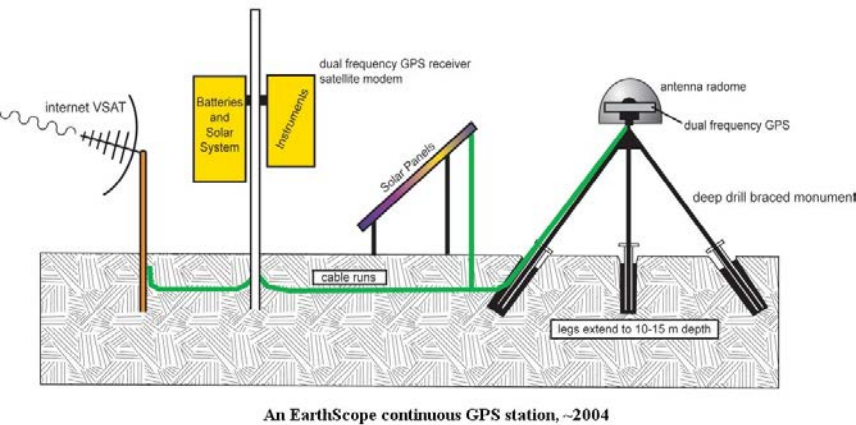
Bird [2007] also showed that a well-constrained fault offset rate requires 4~7 offset features, not just 1 or 2 [his Figure 9].

THEREFORE:

Therefore, all the structures in the Irish Hills, which formed since 5 Ma, should have been studied and modeled to provide geologic constraints on the rates of thrust-faulting.

 I provided one example in Figure 1 of my March 2024 Declaration: Throw of the Obispo Formation at the San Luis Bay-Inferred Coastline thrust fault is 1.6~2.2 km since 5 Ma, implying throw-rate of 0.32~0.44 mm/year, and fault slip rate of 0.76~1.04 mm/year.

If thrusting in the Irish Hills has been symmetrical(?), then a minimum total thrust slip-rate by this method would be 1.52~2.08 mm/year. (However, this neglects any internal blind thrusts.)



#4. GPS geodetic velocities are not useful for site-specific seismic hazard estimation.

PG&E operated a GPS receiver at DCP, and *PG&E* [2015] reported the shortening direction across the Irish Hills ($\sim N15^\circ E$), but not the rate. The *PG&E* [2024] update adds no new geodetic information!

HOWEVER:

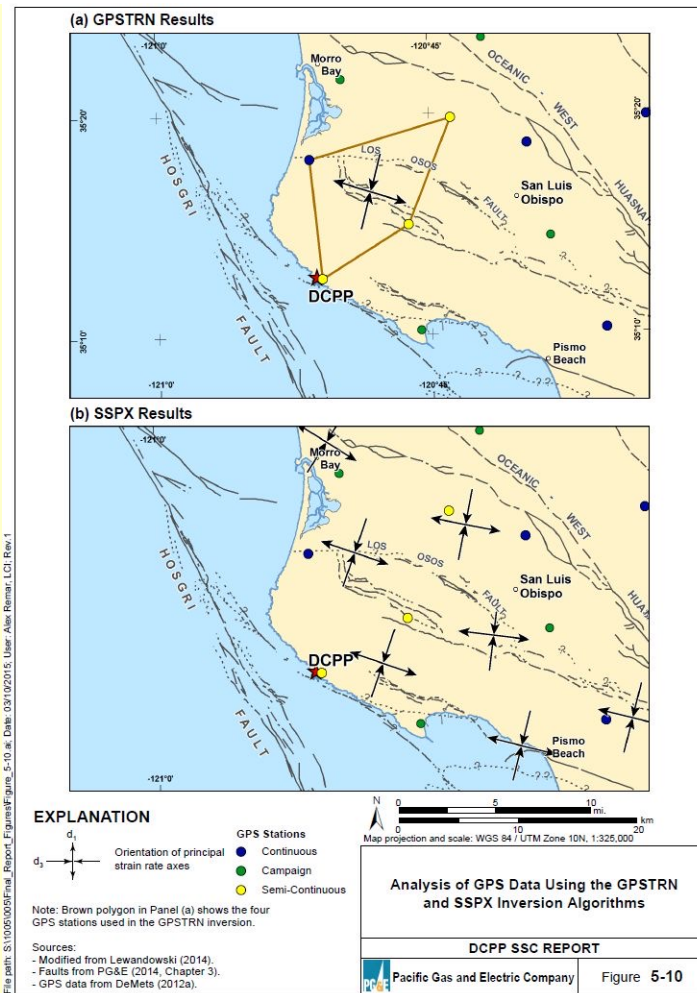
Seismicity has been successfully forecast using only GPS data, both in southern California [*Shen et al.*, 2007] and globally [*Bird et al.*, 2010; *Bird & Kreemer*, 2015]. Therefore, GPS data are useful. Any deformation model used in SSC should fit GPS strain-rate constraints (within their uncertainties).

THEREFORE:

Our two NeoKinema models of neotectonics in the western US [*Field et al.*, 2013; *Shen & Bird*, 2022] had low-resolution F-E grids in the Irish Hills region, but:

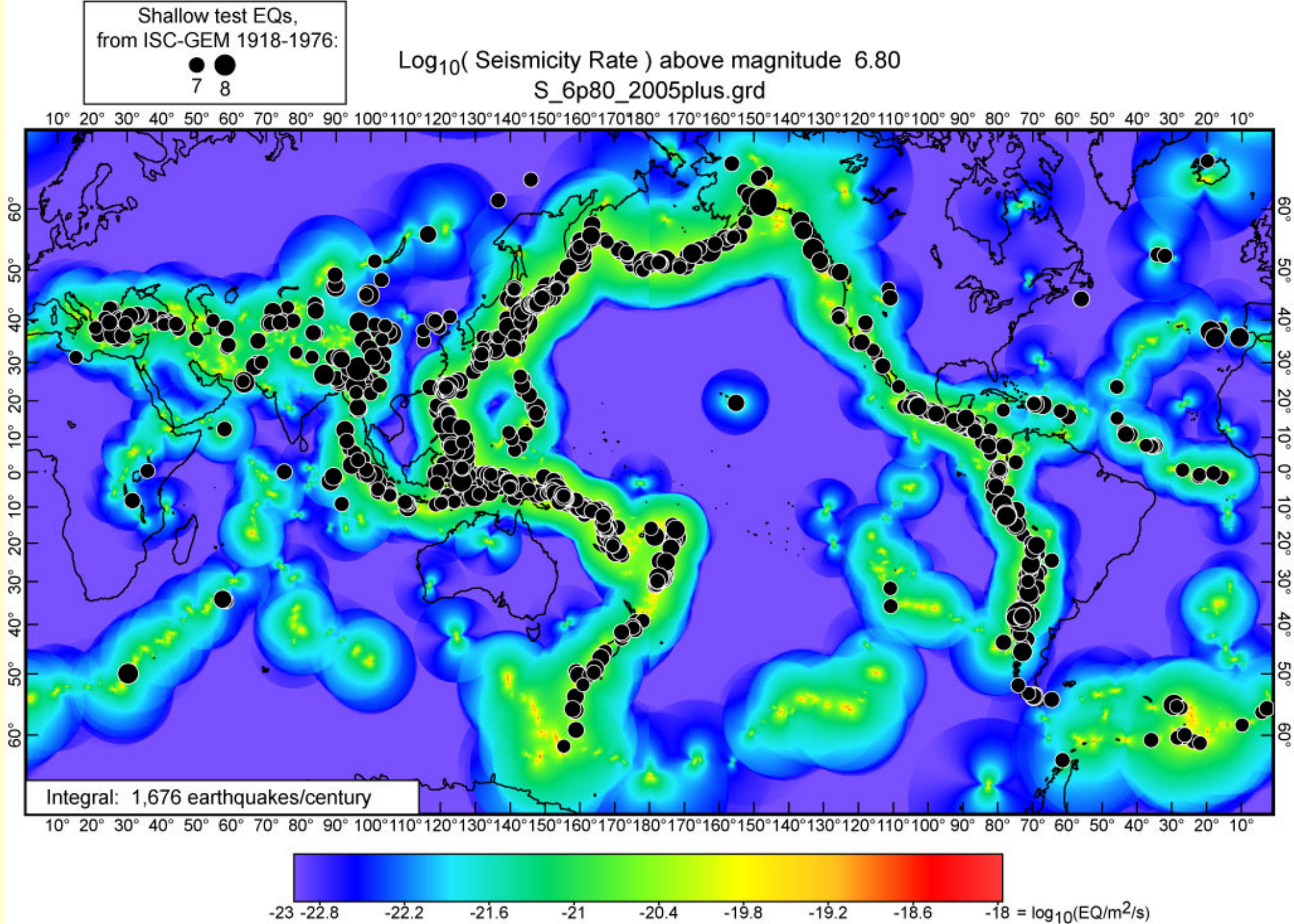
Both showed ~ 2 mm/year shortening across the Irish Hills, implying total thrust fault slip rate of $(\sim 2 \text{ mm/year}) / \cos(25^\circ) = \sim 2.2 \text{ mm/year}$.

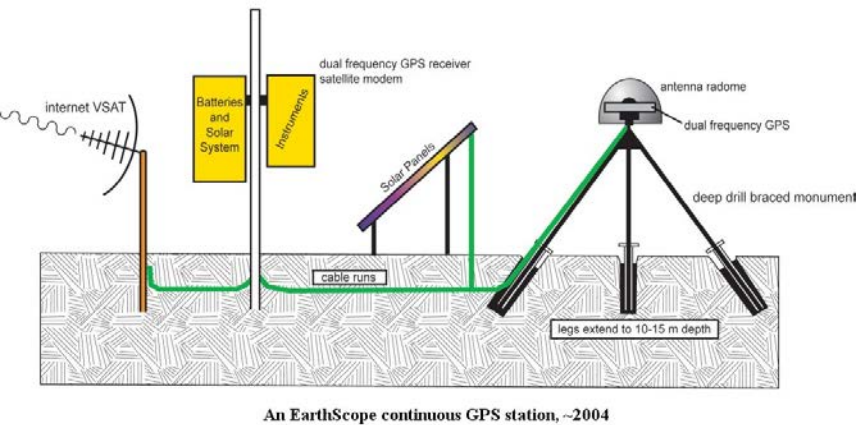
This is the 3rd of 3 independent analytic estimates developed here.



Bird & Kreemer
[2015, BSSA]
global forecast
based on strain-rates
measured by GPS
geodesy.

Historic earthquakes
(black dots) were
not used in creating
this forecast.





#4. GPS geodetic velocities are not useful for site-specific seismic hazard estimation.

PG&E operated a GPS receiver at DCP, and *PG&E* [2015] reported the shortening direction across the Irish Hills ($\sim N15^\circ E$), but not the rate. The *PG&E* [2024] update adds no new geodetic information!

HOWEVER:

Seismicity has been successfully forecast using only GPS data, both in California [*Shen et al.*, 2007] and globally [*Bird et al.*, 2010; *Bird & Kreemer*, 2015].

Therefore, GPS data are useful. Any deformation model used in SSC should fit GPS strain-rate constraints (within their uncertainties).



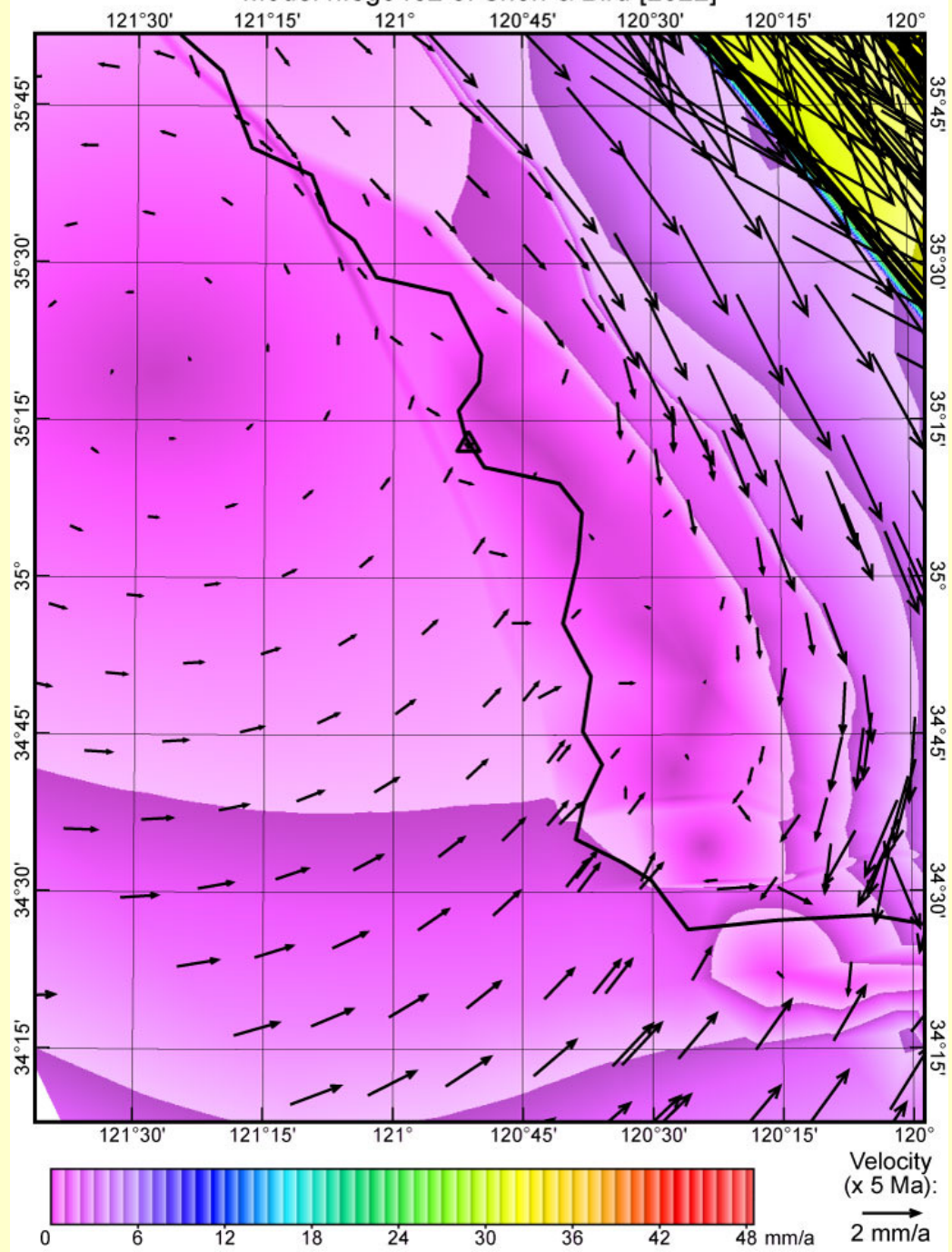
THEREFORE:

Our two NeoKinema models of neotectonics in the western US [*Field et al.*, 2013; *Shen & Bird*, 2022] had low-resolution F-E grids in the Irish Hills region, but:

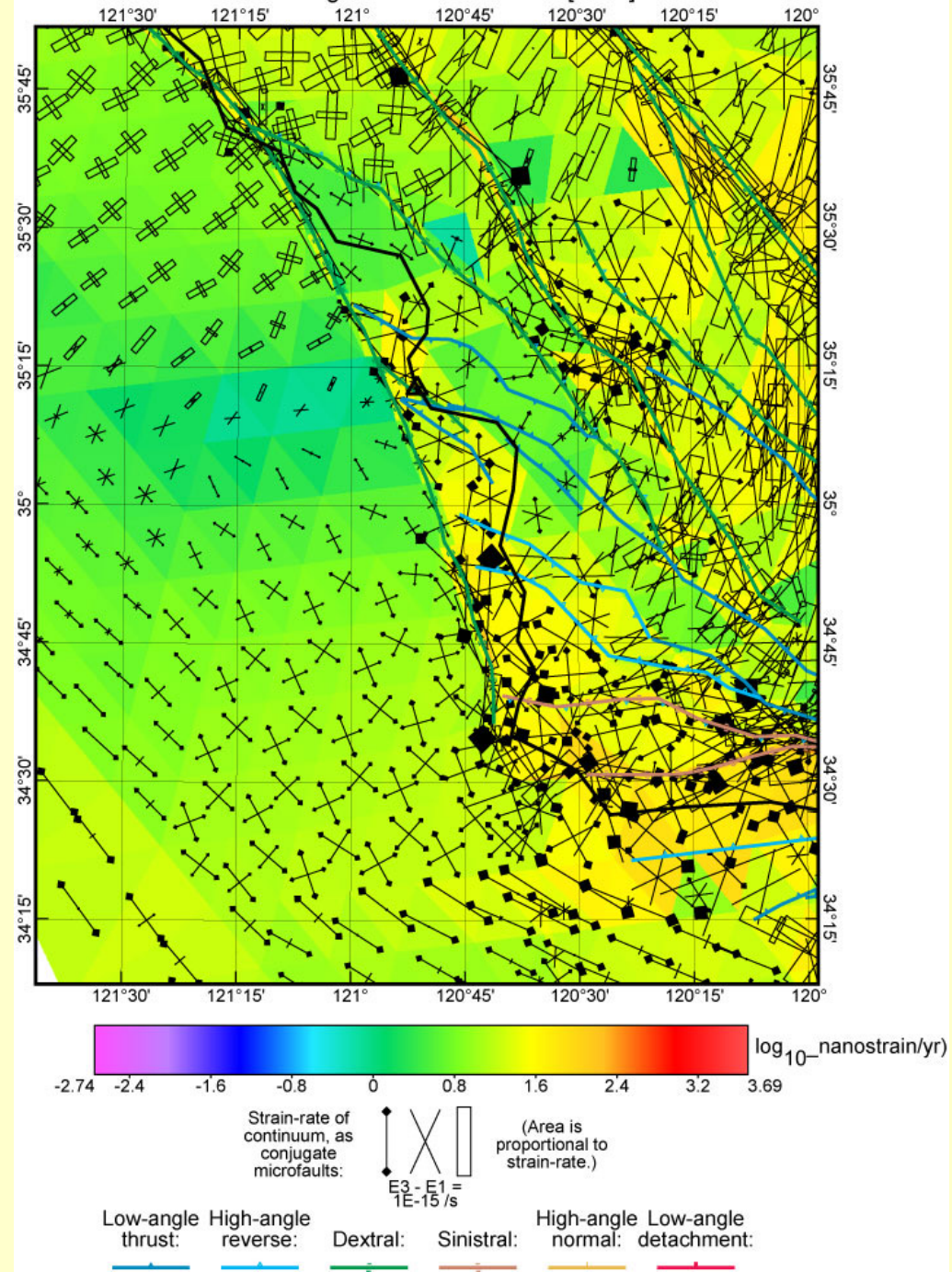
Both showed ~ 2 mm/year shortening across the Irish Hills, implying total thrust fault slip rate of $(\sim 2 \text{ mm/year}) / \cos(25^\circ) = \sim 2.2 \text{ mm/year}$.

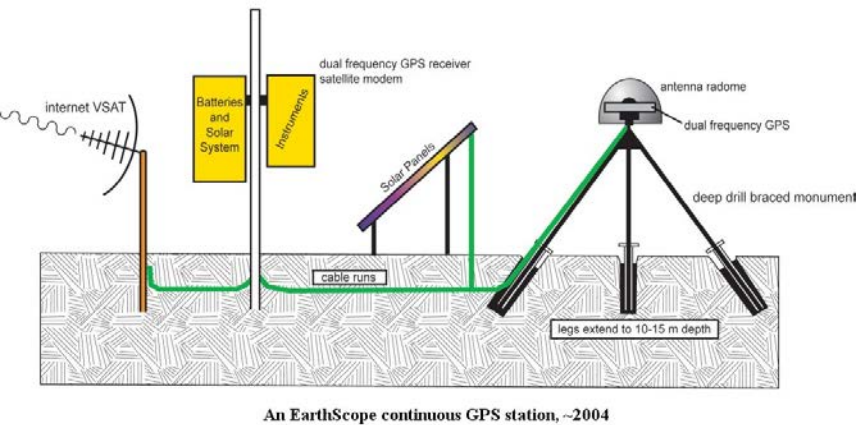
This is the 3rd of 3 independent analytic estimates developed here.

Long-term-average velocity with respect to DCP
Model nfeg0402 of Shen & Bird [2022]



Continuum Strain-Rate Tensors, excluding modeled fault slip
Model nfeg0402 of Shen & Bird [2022]





#4. GPS geodetic velocities are not useful for site-specific seismic hazard estimation.

PG&E operated a GPS receiver at DCP, and *PG&E* [2015] reported the shortening direction across the Irish Hills ($\sim N15^\circ E$), but not the rate. The *PG&E* [2024] update adds no new geodetic information!

HOWEVER:

Seismicity has been successfully forecast using only GPS data, both in California [*Shen*, 2007] and globally [*Bird et al.*, 2010; *Bird & Kreemer*, 2015]. Therefore, GPS data are useful. Any deformation model used in SSC should fit GPS strain-rate constraints (within their uncertainties).

THEREFORE:

Our two NeoKinema models of neotectonics in the western US [*Field et al.*, 2013; *Shen & Bird*, 2022] had low-resolution F-E grids in the Irish Hills region, but:



Both showed ~ 2 mm/year shortening across the Irish Hills, implying total thrust fault slip rate of $(\sim 2 \text{ mm/year}) / \cos(25^\circ) = \sim 2.2 \text{ mm/year}$.

This is the 3rd of 3 independent analytic estimates developed here.

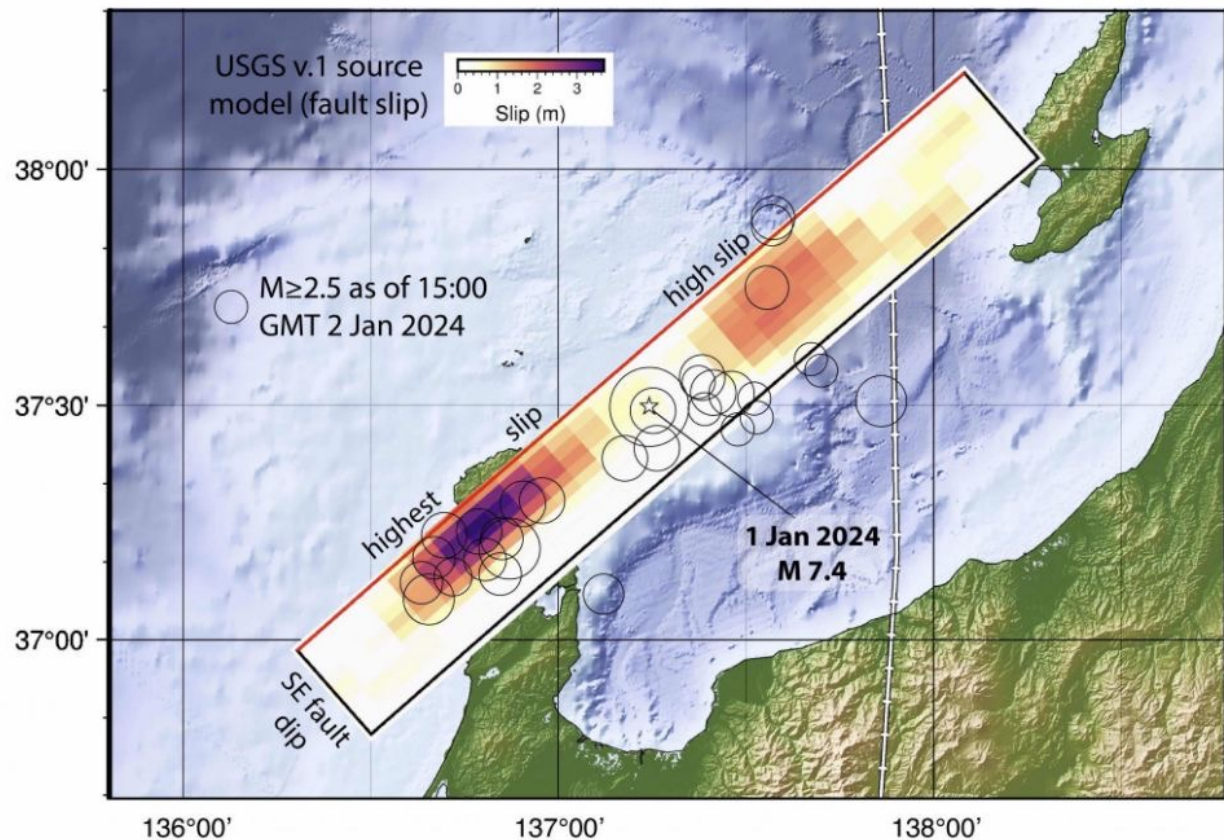
- Each time a **false assumption** was removed, **thrust-faulting activity (seismic potency rate) in the Irish Hills went up by a large factor.**
- It is important to estimate how these factors combine, and how much seismic hazard (and SCDF) is increased at DCP.
- This could be done with a new SSC study and a new SPRA study, except that we cannot afford years of time and millions of \$.
- Instead, we will use a much simpler method to show that the lower limit on seismic hazard (and SCDF) due to thrust-faulting **alone** is much higher than the total hazard claimed by PG&E.
- We will do this by adopting a characteristic great thrust earthquake for this tectonic setting, and then estimating its frequency in the Irish Hills.

A CHARACTERISTIC GREAT THRUST EARTHQUAKE?

The Noto Peninsula on the northwest coast of Japan is tectonically analogous to the Irish Hills: a block of crust now being uplifted between two conjugate intraplate thrust faults.

We learned 2 essential facts from the 2024.01.01 m7.5 earthquake there:

- Mean slip on the seismogenic part of the thrust was 2 m [*USGS finite-fault solution*].
- Peak ground accelerations (PGA) at 5 strong-motion seismometers were 1.0~2.3 g.

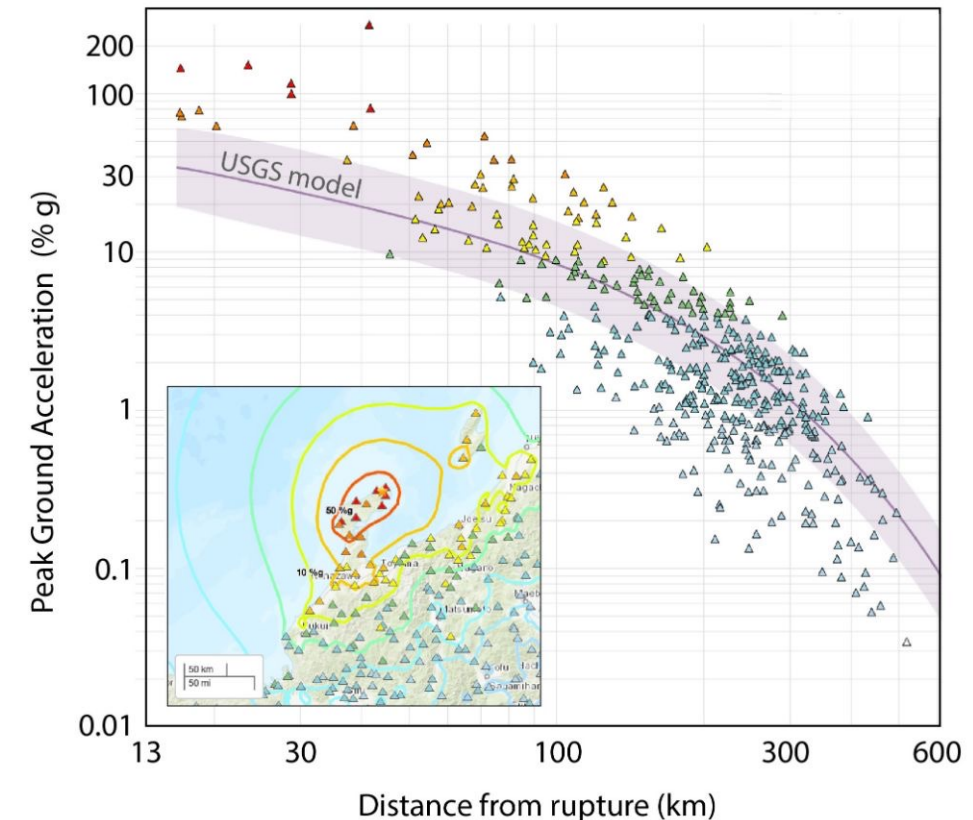


Toda
&
Stein
[2024]

temblor®

Global Risk Solutions Temblor Earth News

predicted the shaking within 40 kilometers of the rupture by factors of two to five.



CONCLUSIONS:

- The two SSC studies by *PG&E* [2015; 2024] seriously underestimated the seismic hazard from thrust-faulting under the Irish Hills because they relied on 4 demonstrably false assumptions.
- Three independent analytic methods give values for the total slip-rate on all shallow-dipping thrust faults under the Irish Hills:
2.8 mm/year, ~2.0 mm/year, or 2.2 mm/year.
- Using the 2024.01.01 Noto Peninsula earthquake as a characteristic great thrust earthquake (with its 2 m of mean slip) yields recurrence times for great thrust earthquakes under the Irish Hills of 715 years, 1000 years, or 910 years, respectively.
- Because such a great thrusting earthquake would cause seismic core damage at DCPP, its seismic core damage frequency (SCDF) is at least

1.4×10^{-3} /year, or 1.0×10^{-3} /year, or 1.1×10^{-3} /year, respectively.

[This is before the hazard contribution from strike-slip faults like the Hosgri is added.]

REFERENCES CITED

- Bird, P. [2007] Uncertainties in long-term geologic offset rates of faults: General principles illustrated with data from California and other western states, *Geosphere*, **3**(6), 577-595; doi:10.1130/GES00127.1, + 9 digital file appendices.
- Bird, P., and C. Kreemer [2015] Revised tectonic forecast of global shallow seismicity based on version 2.1 of the Global Strain Rate Map, *Bull. Seismol. Soc. Am.*, **105**(1), 152-166, doi: 10.1785/0120140129.
- Byerlee, J. [1978] Friction in rocks, *Pure Appl. Geophys.*, **116**, 615-626.
- Field, E. H., G. P. Biasi, P. Bird, T. E. Dawson, K. R. Felzer, D. D. Jackson, K. M. Johnson, T. H. Jordan, C. Madden, A. J. Michael, K. R. Milner, M. T. Page, T. Parsons, P. M. Powers, B. E. Shaw, W. R. Thatcher, R. J. Weldon, II, and Y. Zeng [2013] Unified California Earthquake Rupture Forecast, version 3 (UCERF3)-The time-independent model, *U.S. Geol. Surv. Open-File Rep.*, **2013-1165** (and *Cal. Geol. Surv. Spec. Rep.* **228**, and *Southern California Earthquake Center Pub.* **1792**), 97 pages (main report) + 20 Appendices; <http://pubs.usgs.gov/of/2013/1165/>; including:
- Parsons, T., K. M. Johnson, P. Bird, J. Bormann, T. E. Dawson, E. H. Field, W. C. Hammond, T. A. Herring, R. McCaffrey, Z.-K. Shen, W. R. Thatcher, R. J. Weldon, II, and Y. Zeng [2013] Appendix C—Deformation Models for UCERF3. 66 pages.

Pacific Gas and Electric Company [2014] Geologic Map of the Irish Hills and Adjacent Area, 1:32,000, *DCPP Geologic Mapping Project*, **Ch9.GEO.DCPP.TR.14.01 R0**,
https://www.pge.com/includes/docs/pdfs/safety/systemworks/dcpp/report/Ch9.GEO.DCPP.TR.14.01_R0_Plates.pdf ,
NRC ADAMS Accession No. ML14260A068.

Pacific Gas and Electric Company (PG&E) [2015] Seismic Source Characterization for the Diablo Canyon Power Plant, San Luis Obispo County, California; report on the results of SSHAC level 3 study, Rev. A, March; 652 pages plus Appendices. Available online at <http://www.pge.com/dcpp-ltsp>; downloaded 2023.05.11.

Pacific Gas and Electric Company (PG&E) [2024] Diablo Canyon Updated Seismic Assessment: Response to Senate Bill 846, 1 February 2024, 392 pages.

Shen, Z.-K., and P. Bird [2022] NeoKinema deformation model for the 2023 update to the U.S. National Seismic Hazard Model, *Seismol. Res. Lett.*, **93**, 3037-33052, doi: 10.1785/0220220179.

Shen, Z.-K., D. D. Jackson, and Y. Y. Kagan [2007] Implications of geodetic strain rate for future earthquakes, with a 5-year forecast of M5 earthquakes in southern California, *Seismol. Res. Lett.*, **78**(1), 116-120.

Toda, S., and Stein, R. S. [2024] Intense seismic swarm punctuated by a magnitude 7.5 Japan shock, *Temblor*, <http://doi.org/10.32858/temblor.333>