

Sandia National Laboratories

Albuquerque, New Mexico 87185

March 15, 1986

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Mr. Walton Kelly  
U.S. Nuclear Regulatory Commission  
Mail Stop 623-SS  
Washington, DC 20545

Dear Mr. Kelly:

Enclosed is the monthly report for FIN A-1756, Geochemical Sensitivity Analysis for February 1986.

Please feel free to contact me if you have any questions or comments.

Sincerely,

*Robert M. Cranwell*

R. M. Cranwell  
Supervisor  
Waste Management Systems  
Division 6431

RMC:6431:jm

Enclosure

WM-RES  
WM Record File  
A-1756  
SNL

WM Project 10,11,16  
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PROGRAM: Geochemical Sensitivity  
Analysis

FIN#: A-1756

CONTRACTOR: Sandia National  
Laboratories

BUDGET PERIOD: 10/01/85 -  
9/30/86

DRA PROGRAM MANAGER: W. R. Kelly BUDGET AMOUNT: 265K

CONTRACT PROGRAM MANAGER: R. M. Cranwell FTS PHONE: 844-8368

PRINCIPAL INVESTIGATOR: M. D. Siegel FTS PHONE: 846-5448

#### PROJECT OBJECTIVES

The objective of this project is to provide technical assistance to the NRC in determining the sensitivity of far-field performance assessment calculations to uncertainties in geochemical and hydrological input data and in the representation of geochemical processes in transport models. In Task I, the error in model calculations of integrated radionuclide discharge due to speciation, kinetic and sorption effects will be evaluated. In Task II, the potential importance of organic molecules and colloids will be examined. SNLA will assist the NRC in determining how geochemical processes should be represented in transport models under Task III. Short-term technical assistance will be carried out under Task IV.

#### ACTIVITIES DURING FEBRUARY 1986

##### Task I. Uncertainty in Integrated Radionuclide Discharge

###### • Conceptual Models for Sites

A draft description of the conceptual models of the hydrology of hypothetical basalt and bedded salt sites was written during this month. These descriptions will be a starting point for collaboration with the hydrological data base compilation and hydrological sensitivity analyses being conducted by other NRC contractors (Nuclear Waste Consultants).

A review of previous efforts to rank radionuclides according to their importance in HLW disposal was initiated during February. The result of this review will be a letter report which will later be included in the introductory sections of the final report for this project.

- Solubility/Speciation Effects

Data for minerals found in basaltic geological settings were added to the Aqueous Solutions Data Base. In addition, data on aqueous species were updated and corrected during this month. A draft report summarizing the status of the ASD and potential uses in sensitivity analyses is in preparation.

Malcolm Siegel attended the USGS computer modeling course "Geochemistry for Ground-Water Systems," in Denver, Colorado on February 26 to March 7, 1986.

- Sorption Effects

A complete first draft of a report describing the application of the Stanford Generalized Model for Adsorption to radionuclide sorption was received during February. The document is under review at SNLA and will be sent to the NRC after necessary revisions are made. A preliminary list has been compiled of geochemical systems for which the effects of fluid composition on theoretical  $K_d$ 's can be modeled. (See Attachment 1.) The systems which are most relevant to HLW disposal will be selected for modeling at SNLA and Stanford University.

- Kinetic and Dynamic Effects

The final version of the paper "Approximate Methods to Calculate Radionuclide Discharges for Performance Assessment of HLW Repositories in Fractured Rock" was completed and is appended as Attachment 2.

A user's manual, user-friendly interface, and a well-documented source code for the chemical transport simulator TRANQL are nearly complete. It is anticipated that the code will be transferred to SNLA during March.

### Other Activities

A revised Schedule 189 for this project was prepared during February and approved by the NRC. Under this new Schedule, a new task, "Transfer of Sensitivity Analysis Tools," has been added to the Geochemical Sensitivity Analysis project.

### Trips

- M. Siegel (6431) attended the U.S. Geological Survey course "Geochemistry for Ground-Water Systems" at the USGS National Training Center in Denver, Colorado on February 24 - March 7, 1986. The course dealt with computer models developed by the USGS to model the chemistry of ground waters. The program emphasized four major topics: 1) thermodynamic basis for

geochemical models, 2) major rock/water reactions, 3) use of the USGS codes WATEQF, BALANCE and PHREEQE, and 4) case studies illustrating the application of the codes to ground-water systems studied by the USGS.

#### Allocation of Resources

Task I. 100%

# ATTACHMENT I

## REPORT ON GEOCHEMICAL SCENARIOS

<u>ADSORBATE<sup>+</sup></u>	<u>ADSORBENTS<sup>*</sup></u>	<u>SOLUTION VARIABLES<sup>#</sup></u>
UO <sub>2</sub> <sup>2+</sup>	$\alpha$ -FeOOH, am-Fe(OH) <sub>3</sub>	P <sub>CO<sub>2</sub></sub> , Ca <sup>2+</sup> , EDTA, pH
Pu	$\alpha$ -FeOOH	P <sub>CO<sub>2</sub></sub> , Ca <sup>2+</sup> , EDTA, pH
Np	am-Fe(OH) <sub>3</sub>	P <sub>CO<sub>2</sub></sub> , Ca <sup>2+</sup> , EDTA, pH
Ni <sup>2+</sup>	SiO <sub>2</sub> , $\alpha$ -FeOOH, am-Fe(OH) <sub>3</sub> , $\alpha$ -Al <sub>2</sub> O <sub>3</sub>	EDTA, pH
SeO <sub>3</sub> /SeO <sub>4</sub>	$\alpha$ -FeOOH, am-Fe(OH) <sub>3</sub> , $\alpha$ -Al <sub>2</sub> O <sub>3</sub>	pH
Pb <sup>2+</sup>	$\alpha$ -FeOOH, am-Fe(OH) <sub>3</sub> , $\alpha$ -Al <sub>2</sub> O <sub>3</sub> , SiO <sub>2</sub>	pH, EDTA, P <sub>CO<sub>2</sub></sub> , Ca <sup>2+</sup>
Cd <sup>2+</sup>	am-Fe(OH) <sub>3</sub> , $\alpha$ -Al <sub>2</sub> O <sub>3</sub> , TiO <sub>2</sub> , SiO <sub>2</sub>	P <sub>CO<sub>2</sub></sub> , Ca <sup>2+</sup> , Cl <sup>-</sup> , EDTA, pH
Zn <sup>2+</sup>	$\alpha$ -FeOOH, am-Fe(OH) <sub>3</sub> , $\alpha$ -Al <sub>2</sub> O <sub>3</sub> , SiO <sub>2</sub>	P <sub>CO<sub>2</sub></sub> , Ca <sup>2+</sup> , EDTA, pH
Cu <sup>2+</sup>	$\alpha$ -FeOOH, am-Fe(OH) <sub>3</sub> , $\alpha$ -Al <sub>2</sub> O <sub>3</sub> , TiO <sub>2</sub> , SiO <sub>2</sub>	P <sub>CO<sub>2</sub></sub> , Ca <sup>2+</sup> , EDTA, pH

\* In every case where more than one solid (adsorbent) is available, it is possible to model transitions from one solid to another. Total surface area can be varied.

# P<sub>CO<sub>2</sub></sub> is intended to represent several different possible modes of representing total carbonate in the system.

+ The total concentration of adsorbate can be varied in the simulations if needed.



## ATTACHMENT II

### APPROXIMATE METHODS TO CALCULATE RADIONUCLIDE DISCHARGES FOR PERFORMANCE ASSESSMENT OF HLW REPOSITORIES IN FRACTURED ROCK

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#### ABSTRACT

Three approximate methods appear useful for calculating radionuclide discharges in fractured, porous rock: (1) a semi-infinite-medium approximation where radionuclide diffusion rates into the matrix are calculated assuming a semi-infinite matrix; (2) a linear-driving-force approximation where radionuclide diffusion rates into the matrix are assumed to be proportional to the difference between bulk concentrations in the fracture fluid and in the matrix pore water; and (3) an equivalent-porous-medium approximation where radionuclide diffusion rates into the matrix are calculated assuming that the time rate of change of the bulk radionuclide concentration in the matrix is proportional to the time rate of change of the radionuclide concentration in the fracture fluid. A preliminary evaluation of these approximations was made by considering transport of a single radionuclide in saturated, porous rock containing uniform, parallel fractures. It was assumed that fluid flow was one-dimensional, the nuclide existed as a single chemical species, and radioactive decay and production of the nuclide were negligible. Criteria for application of each approximation were derived in terms of fundamental physicochemical parameters. For parameter values satisfying each of the criteria, the respective errors in radionuclide discharges calculated using the approximations were examined by comparing those discharges with discharges calculated rigorously. In addition, discharges were calculated with the computer code NWFT/DVM, which was developed at Sandia National Laboratories for use in performance assessment calculations. Agreement among results calculated with the analytical exact solution, the analytical linear-driving-force approximation and the numerical linear-driving-force approximation was good for a variety of hydrological conditions. The applicability of each approximation to performance assessment for repositories in basalt, granite, and tuff was shown using site-specific hydrologic and geochemical parameters.

#### INTRODUCTION

Performance assessment requires calculating radionuclide discharges for many sets of conditions for each of several scenarios. General use of rigorous convective-diffusive transport models would be impractical for performance assessment of HLW repositories in fractured, porous rock. While such rigorous calculations are desirable for demonstrating detailed understanding of physicochemical phenomena, they would be unnecessary for risk assessment if upper bounds for radionuclide discharges could be obtained from approximate models. Three approximate methods for calculating radionuclide discharges in fractured, porous rock can be used to minimize the number of rigorous computations: (1) a semi-infinite-medium approximation where radionuclide diffusion rates into the matrix are calculated assuming a semi-infinite matrix; (2) a linear-driving-force approximation where radionuclide diffusion rates into the matrix are assumed to be proportional to the difference between bulk concentrations in the fracture fluid and in the matrix pore water; and (3) an equivalent-porous-medium approximation where radionuclide diffusion rates into the matrix are calculated assuming that the time rate of change of the bulk radionuclide concentration in the matrix is proportional to the time rate of change of the radionuclide concentration in the fracture fluid.

A preliminary evaluation of each approximation was made by considering a relatively simple system involving transport of a single radionuclide in saturated, porous rock containing uniform, parallel

fractures. It was assumed that fluid flow was one-dimensional, the nuclide existed as a single chemical species, and radioactive decay and production of the nuclide were negligible. In the discussion below, the rigorous transport equations for that system are described. Each approximation is discussed; the corresponding transport equations are given, and criteria for application of each approximation are derived in terms of fundamental physical and chemical parameters which are amenable to measurement in the laboratory or field. For parameter values satisfying each of the criteria, the respective errors in radionuclide discharges calculated using the approximations are examined by comparing those discharges with discharges calculated rigorously. The applicability of each approximation to performance assessment for repositories in basalt, granite, and tuff is discussed using site-specific hydrologic and geochemical parameters.

#### THEORY

##### Radionuclide Transport in Fractured, Porous Rock

Consider a region of saturated, porous rock containing a system of uniform, parallel fractures which divide the porous matrix into parallel flat plates (Fig. 1). Assume that: (1) fluid flow occurs only in the x-direction and is negligible in the porous matrix; (2) effects due to hydrodynamic dispersion are negligible; (3) radionuclide concentrations in the fracture fluid are uniform across the fracture cross section; (4) local chemical equilibrium exists at the fracture-matrix and pore-water-matrix inter-

faces; (5) bulk radionuclide diffusion in the pore water occurs only in the z-direction; (6) surface diffusion of nuclides in the interfacial regions between pore water and mineral phases negligibly affects transport in the porous matrix; (7) radionuclide sorption is reversible and can be represented by linear isotherms; (8) colloidal transport of nuclides is negligible; (9) radionuclides exist as a single chemical species; (10) radioactive decay and production of the nuclide of interest are negligible; and (11) diffusion coefficient is a constant. Then, the material balance for a radionuclide in the fracture fluid is

$$\frac{\partial C_f}{\partial t} + v \frac{\partial C_f}{\partial x} + m_f \frac{\partial M}{\partial t} = 0 \quad (1)$$

and in the porous matrix

$$\frac{\partial C_m}{\partial t} = D_e \frac{\partial^2 C_m}{\partial z^2} \quad (2)$$

The various terms in Eqs. (1) and (2) are defined in Table I. Appropriate initial and boundary conditions for Eq. (1) are  $C_f(x,0) = 0$ ,  $M(x,0) = 0$ , and  $C_f(0,t) = C_0 = \text{a constant}$ . Appropriate conditions for Eq. (2) are  $C_m(x,z,0) = 0$ ,  $C_m(x,B,t) = C_f(x,t)$ , and  $\partial C_m(x,0,t)/\partial z = 0$ . The solution to Eqs. (1) and (2) is obtained analogously to Rosen's<sup>1</sup> results for packed beds of spheres. The fracture fluid concentration  $C_f(x,t)$  is given by

$$\frac{C_f(x,t)}{C_0} = \frac{1}{2} + \frac{2}{\pi} \int_0^{\infty} e^{-\bar{x}H_{D1}(B)} \sin\left[\bar{y}B^2 - \bar{x}H_{D2}(B)\right] \frac{dB}{B} \quad (3)$$

where

$$\bar{x} = \frac{D_e \phi_m R_m m_f x}{v B^2} = \left(\frac{D}{\alpha}\right) \phi_m m_f x \quad (4)$$

$$\bar{y} = \frac{2D_e(t - \frac{x}{v})}{B^2}$$

$$H_{D1}(B) = \frac{B[\sinh(2B) - \sin(2B)]}{\cosh(2B) + \cos(2B)}$$

$$H_{D2}(B) = \frac{B[\sinh(2B) + \sin(2B)]}{\cosh(2B) + \cos(2B)}$$

$B$  = variable of integration.

The integral in Eq. (3) must be evaluated numerically. However, for large  $\bar{x}$ , say  $\bar{x} \geq 50$ ,  $C_f(x,t)$  can be approximated by:

$$\frac{C_f(x,t)}{C_0} = \frac{1}{2} \left\{ 1 + \operatorname{erf} \left[ \frac{\bar{y}}{2\sqrt{\frac{1}{3\bar{x}}}} - 1 \right] \right\} \quad (4)$$

Eq. (4) is obtained from Eq. (3) using the identity<sup>2</sup>

$$\int_0^{\infty} e^{-B^2} \sin(2\bar{y}B) \frac{dB}{B} = \frac{\pi}{2} \operatorname{erf}(\bar{y})$$

and noting that  $H_{D1} \rightarrow (4/3)B^4$  and  $H_{D2} \rightarrow 2B^2$  as  $B \rightarrow 0$ .

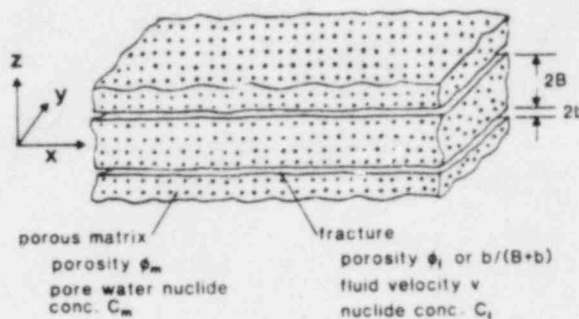


Fig. 1. Schematic representation of fractured, porous rock. Origin of coordinate system is at center of block of porous rock.

#### Approximations

**General.** The rigorous solution to Eqs. (1) and (2) can be obtained using Duhamel's theorem to express the term  $M(x,t)$  in Eq. (1) as

$$M(x,t) = \frac{\phi_m R_m}{B} \int_0^B \int_0^t C_f(x,t') \frac{\partial H(z,t-t')}{\partial t} dt' dz \quad (5)$$

where  $H(z,t)$  is the solution to Eq. (2) when the surface boundary condition is replaced by  $C_m(x,B,t) = 1$ . In each of the three approximations, a simpler expression for  $M(x,t)$ , or  $\partial M(x,t)/\partial t$ , is substituted for Eq. (5), or the corresponding expression for  $\partial M(x,t)/\partial t$ , so that the resulting solution for  $C_f(x,t)$  is much simpler to obtain than Eq. (3). For convenience below, let  $\bar{H}(t)$  denote the volume-averaged nuclide concentration in pore water when surface concentration is unity (see Table I), and write Eq. (5) as

$$M(x,t) = \phi_m R_m \int_0^t C_f(x,t') \frac{\partial \bar{H}(t-t')}{\partial t} dt' \quad (6)$$

**Semi Infinite Medium.** In this approximation, the basic assumption is that the porous matrix is so large that the concentration  $C_m(x,0,t)$  of the diffusing nuclides can be considered negligible during the time interval of interest. The term  $\partial \bar{H}/\partial t$  in Eq. (6) is then obtained from the solution to the diffusion equation for semi-infinite, flat plates having unit surface concentration. For fluid flowing between parallel plates separated by aperture  $2b$ , the expression for  $C_f(x,t)$  is given by<sup>3</sup>

$$\frac{C_f(x,t)}{C_0} = \operatorname{erfc} \left( \frac{\phi_m R_m x \sqrt{D_e}}{2bv \sqrt{t - \frac{x}{v}}} \right) = \operatorname{erfc} \left( \frac{\sqrt{2}}{2} \sqrt{\frac{\bar{x}}{\bar{y}}} \right) \quad (7)$$



Table 1  
Nomenclature

Term	Definition
B	half thickness of matrix between fractures
$\bar{B}$	$B+b$ , half spacing between fractures
b	half of the fracture aperture
$C_f$	radionuclide concentration in the fracture fluid
$C_m$	radionuclide concentration in the matrix pore water
D	molecular diffusion coefficient for the radionuclide in the matrix pore water
$D_e$	$D/\alpha^2 R_m$
$H(z,t)$	solution to Eq. 2 when the surface boundary condition is replaced by $C_m(x, \delta, t) = 1$
$\bar{H}(t)$	$\frac{1}{B} \int_0^B H(z,t) dz$ , the volume average of $H(z,t)$
$\phi_f$	$b/\bar{B}$ , porosity associated with the fractures
$\phi_m$	matrix porosity
$m_f$	$(1-\phi_f)/\phi_f$
$m_m$	$(1-\phi_m)/\phi_m$
M	$\phi_m R_m \left( \frac{1}{B} \right) \int_0^B C_m dz$ , the average matrix radionuclide concentration
$R_m$	$(1+m_m \Gamma_0)$ , radionuclide retardation factor for the porous matrix
t	time
v	mass-averaged fracture fluid velocity
x	spatial coordinate in the direction of the bulk fluid motion
z	spatial coordinate in the direction of diffusion in the porous matrix
$\epsilon^2$	tortuosity/constrictivity factor for the porous matrix
$\Gamma_0$	slope of the linear portion of the dimensionless sorption isotherm (fluid- and solid-phase concentrations both expressed as mass or moles per unit volume)
$\bar{x}$	$\frac{D m_f \phi_m x}{\alpha^2 v B^2}$
$\bar{y}$	$\frac{2D_e(t-x/v)}{B^2}$

The relative errors in  $C_f$  which result from using Eq. (7) rather than Eq. (3) are determined by the corresponding errors in the value for  $M(x,t)$ , which in turn depend on the errors in the value of  $\partial \bar{H}/\partial t$ . Criteria for using Eq. (7) can be developed as follows. First, consider one-dimensional diffusion into semi-infinite flat plates such as depicted in Fig. 1. Let  $z' = B - z$ , and  $C_m(z' = 0, t) = 1$ . The corresponding expressions for  $\bar{H}(z', t)$  and  $H(t)$  are given, respectively, by<sup>4</sup>

$$H(z', t) = \text{erfc} \left( \frac{z'}{2\sqrt{D_e t}} \right) \quad (8)$$

and

$$\bar{H}(t) = \frac{2}{\sqrt{\pi}} \left( \frac{D_e t}{B^2} \right)^{1/2} \quad (9)$$

As the time interval of interest increases, the value of  $\bar{H}(t)$  determined from Eq. (9) will be increasingly in error, as can be seen by inspection of the expression from the exact solution<sup>5</sup> for "short times"

$$\bar{H}(t) = 2 \left( \frac{D_e t}{B^2} \right)^{1/2} \left\{ \frac{1}{\sqrt{\pi}} + 2 \sum_{n=1}^{\infty} (-1)^n \text{ierfc} \left( \frac{nB}{\sqrt{D_e t}} \right) \right\} \quad (10)$$

Let  $\bar{H}_a$  and  $\bar{H}_e$  denote the values of  $\bar{H}$  obtained from Eqs. (9) and (10), respectively, and let  $F_a = (\partial \bar{H}_a / \partial t) / (\partial \bar{H}_e / \partial t)$ . Eq. (6) then can be written as

$$M(x, t) = \phi_m R_m \int_0^t C_f(x, t') \left[ \frac{1}{F_a(t-t')} \right] \frac{\partial \bar{H}_a(t-t')}{\partial t} dt' \quad (11a)$$

and by the mean value theorem<sup>6</sup>

$$M(x, t) = \frac{\phi_m R_m}{F_a(t-\eta)} \int_0^t C_f(x, t') \frac{\partial \bar{H}_a(t-t')}{\partial t} dt' \quad (11b)$$

where  $0 \leq \eta \leq t$ . The term  $F_a(t-\eta) = F_a^*$  represents some mean value of  $F_a(t-t')$ . By definition  $F_a \geq 1$ .

When  $F_a = 1$ ,

$$M(x, t) = \phi_m R_m \int_0^t C_f(x, t') \frac{\partial \bar{H}_a(t-t')}{\partial t} dt' \quad (12)$$

and when  $F_a > 1$ , the ratio of approximate to exact values of  $M(x, t)$  is given by  $F_a^*$ . Since  $\bar{H}_a$ ,  $\bar{H}_e$ ,  $\partial \bar{H}_a / \partial t$ ,  $\partial \bar{H}_e / \partial t$ , and  $F_a$  are single-valued functions of  $t$ , a one-to-one relationship exists between  $F_a$  and  $\bar{H}_e$ . Values of  $F_a$  as a function of  $\bar{H}_e$  are shown in

Fig. 2. For  $\bar{H}_e < 0.5$ ,  $F_a \approx 1$ , while for  $\bar{H}_e > 0.5$ ,  $F_a$  increases monotonically and becomes large for  $\bar{H}_e > 0.8$ . The value of  $F_a^*$  will be between zero and the value of  $F_a(\bar{H}_e)$  corresponding to time  $t$ . For purposes here, it appears reasonable to estimate  $F_a^*$  by the average  $\bar{F}_a$  of  $F_a$  with respect to  $\bar{H}_e$ , that is

$$F_a^*(\bar{H}_e) = \bar{F}_a(\bar{H}_e) = \frac{1}{\bar{H}_e} \int_0^{\bar{H}_e} F_a(\bar{H}) d\bar{H} \quad (13)$$

Values of  $\bar{F}_a$  also are shown in Fig. 2. When  $\bar{H}_e \leq 0.5$ , the values of both  $F_a$  and  $\bar{F}_a$  are about 1.0. When  $0.5 < \bar{H}_e \leq 0.7$ , the value of  $\bar{F}_a$  is about 1.0, and  $F_a$  is 1.2 or less. When  $0.7 < \bar{H}_e \leq 0.92$ , the value of  $\bar{F}_a$  is between about 1.0 and 1.2, and  $F_a$  is between 1.2 and 2.8. When  $\bar{H}_e > 0.92$ ,  $F_a$  and especially  $\bar{F}_a$  become large. Therefore, when  $\bar{H}_e \leq 0.5$ , the ratio of approximate to exact values of  $M(x,t)$  from Eqs. (12) and (6), respectively, would be about 1.0. When  $0.5 < \bar{H}_e \leq 0.7$ , the ratio would be about 1.0 and would not be greater than 1.2. When  $0.7 < \bar{H}_e \leq 0.92$ , the ratio would be about 1.2 or less and would not be greater than 2.8.

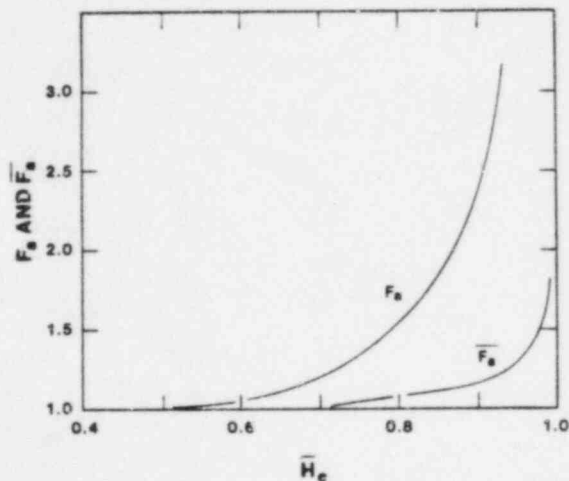


Fig. 2. Values of  $F_a$  and  $\bar{F}_a$  versus  $\bar{H}_e$ .

Criteria for applying the semi-infinite-medium approximation, that is, calculating  $M(x,t)$  from Eqs. (12) and (9) rather than Eqs. (6) and (10), will depend on the errors which are acceptable in  $M(x,t)$  and in calculated radionuclide discharges. When solving Eq. (1) with either approximate or exact expressions for  $M(x,t)$ , those expressions actually appear as  $M(x,t - x/v)$ , since  $C_f(x,t) = 0$  if  $t < x/v$  and only convective transport occurs in the  $x$ -direction in the fractures. To examine errors between approximate and exact solutions, the value for  $\bar{H}$  should correspond to the contact time  $t^* = t - x/v$  rather than the elapsed time  $t$ . Conditions under which the semi-infinite-medium approximation are useful could be defined in terms of  $t^*$ . However, the use of maximum mean radionuclide contact time would lead to criteria which are more restrictive and conservative.

Let  $\theta_m$  denote the maximum mean radionuclide residence time along the flow path of interest. In which case,  $\theta_m$  corresponds to the fracture fluid and pore water being in equilibrium with respect to radionuclide diffusion. As discussed later,  $\theta_m$  is given by  $\theta_m = (x/v)R_f$ , where  $x$  corresponds to the

flow path length, and  $R_f$  is the radionuclide retardation factor for the fracture fluid and is given by  $R_f = 1 + \rho_f \phi_m R_m$ . By analogy with  $t^*$ , define the contact residence time  $\theta_m^*$  as  $\theta_m - x/v$ .

First, consider cases where  $t^* < \theta_m^*$ . From Table II,  $\bar{H}_e \leq 0.5$  when  $D_e t^*/B^2 \leq 0.2$ ,  $\bar{H}_e \leq 0.7$  when  $D_e t^*/B^2 \leq 0.4$ , and  $\bar{H}_e \leq 0.92$  when  $D_e t^*/B^2 \leq 1.0$ . If  $\theta_m \leq 0.2B^2/D_e$ , then  $\theta_m^* \leq 0.2B^2/D_e$  for all  $x$ , and perturbations in radionuclide concentrations would be exposed to an essentially semi-infinite porous matrix for which  $F_a \approx 1.0 \approx \bar{F}_a$  throughout. Radionuclide discharges then could be obtained from Eq. (7). If  $0.2 < \theta_m(B^2/D_e) \leq 0.4$ , perturbations in radionuclide concentrations would be exposed to a porous matrix for which  $F_a \leq 1.2$  and  $\bar{F}_a \approx 1.0$ . Again, radionuclide discharges could be obtained from Eq. (7); however, as  $t$  and  $\theta_m$  approach  $0.4B^2/D_e$ , small relative errors in  $M(x,t)$  and, therefore, in values of  $C_f/C_0$  should be expected since  $F_a$  will be greater than 1.0 for some values of  $x$ . If  $0.4 < \theta_m(B^2/D_e) \leq 1.0$ , perturbations in radionuclide concentrations would be exposed to a porous matrix for which  $F_a \leq 2.8$  and  $\bar{F}_a \leq 1.2$ . The relative errors in radionuclide discharges obtained from Eq. (7) would correspond to relative errors in the value of  $M(x,t)$  which are of the order of 20 percent as  $t$  and  $\theta_m$  approach  $1.0B^2/D_e$ .

Table II

Values of  $\bar{H}_e(t)$  Versus  $D_e t/B^2$

$D_e t/B^2$	$\bar{H}_e(t)$	$D_e t/B^2$	$\bar{H}_e(t)$
0.0001	0.01	0.09	0.34
0.001	0.04	0.10	0.36
0.003	0.06	0.20	0.50
0.005	0.08	0.30	0.61
0.008	0.10	0.40	0.70
0.01	0.11	0.50	0.76
0.02	0.16	0.60	0.81
0.03	0.20	0.70	0.86
0.04	0.23	0.80	0.89
0.05	0.25	0.90	0.91
0.06	0.28	1.00	0.93
0.07	0.30	1.50	0.98
0.08	0.32	2.00	0.99

Next, consider cases where  $t^* > \theta_m^*$ . For a given ratio of  $\theta_m/(B^2/D_e)$ , values of  $\bar{F}_a$  will be larger than when  $t^* < \theta_m^*$ . Relative errors in  $M(x,t)$ , and the corresponding errors in radionuclide discharges, also should be larger. However, the radionuclide flux,  $-D_e \partial H(0,t)/\partial x$ , is proportional to  $1/\sqrt{t}$ , which follows from Eq. (8). Therefore, while the relative error associated with the semi-infinite-medium approximation increases, the relative amount of the radionuclide diffusing into the porous matrix decreases, thereby reducing the overall effect that those errors have on radionuclide discharges. In examples given later, it is shown that when  $t^* > \theta_m^*$ , the relative errors in radionuclide discharges do not increase appreciably beyond those corresponding to  $t^* = \theta_m^*$ .

If only small relative errors in calculated radionuclide discharges are acceptable, the criterion for applying the semi-infinite medium approximation is  $\theta_m = (x/v)R_f \leq 0.2B^2/D_e$  or,

$$\frac{D_e R_f x}{v B^2} \leq 0.2$$

Often,  $\phi_f \ll 1$ , and  $\phi_m > \phi_f$ . Under those conditions,  $R_f = m_f \phi_m R_m$  since  $R_m \geq 1$ . Then, the criterion for applying the approximation can be written as

$$\bar{\alpha} = \frac{D_e m_f \phi_m R_m x}{v B^2} = \left(\frac{D}{2}\right) m_f \phi_m x \leq 0.2 \quad (14a)$$

when only small relative errors are acceptable.

However, it seems reasonable to define a less restrictive criterion. Parameter values in the expression for  $M(x,t)$  may often involve uncertainties of 20 to 30 percent or greater, which result from inherent variations in the physical and chemical properties of geomedia. If the criterion  $\theta_m = (x/v) R_f \leq B^2/D_e$  is used, relative errors in  $M(x,t)$  would be on the order of 20 percent for  $t \leq \theta_m$ , and are similar to, or less than, possible uncertainties in parameter values. Hence

$$\bar{\alpha} = \frac{D_e m_f \phi_m R_m x}{v B^2} = \left(\frac{D}{2}\right) m_f \phi_m x \leq 1 \quad (14b)$$

when relative errors of about 20 percent in  $M(x,t)$  and the corresponding errors in calculated discharges are acceptable.

Linear Driving Force. In the usual form of this approximation, the radionuclide flux into the porous matrix is assumed to be proportional to the difference between the surface and average matrix concentrations, and the approximation expressed as

$$\frac{dM}{dt} = k a_m (\phi_m R_m C_f - M) \quad (15)$$

where  $a_m$  is the surface area of the porous matrix contacting the fracture fluid per unit volume of matrix, and  $k$  is a constant mass transfer coefficient which is evaluated analytically below. When Eqs. (1) and (15) are solved, the resulting expression for  $C_f(x,t)$  is

$$\frac{C_f(x,t)}{C_0} = J(n,m) = 1 - e^{-m} \int_0^n e^{-B} I_0(2\sqrt{B}) dB \quad (16)$$

where  $a_m = \frac{1}{B}$

$$m = \frac{k}{B} \left( t - \frac{x}{v} \right)$$

$$n = \frac{k \phi_m R_m x}{B v}$$

and  $B$  is again a variable of integration.<sup>7</sup> Values of the function  $J(n,m)$  have been tabulated extensively,

and for  $nm > 3600$  the function is given approximately by

$$J(n,m) = \frac{1}{2} \operatorname{erfc}(\sqrt{n} - \sqrt{m}) \quad (17)$$

In these analyses, the term  $\partial \bar{H}/\partial t$  in Eq. (6) was approximated using a linear-driving-force expression for diffusion into flat plates having unit surface concentration, that is

$$\frac{d\bar{H}}{dt} = k a_m (1 - \bar{H}) \quad (18)$$

For  $\bar{H}(0) = 0$ ,

$$\bar{H} = 1 - \exp(-k a_m t) \quad (19)$$

and

$$\frac{d\bar{H}}{dt} = k a_m \exp(-k a_m t) \quad (20)$$

Substituting Eq. (20) into Eq. (6) gives

$$M(x,t) = \phi_m R_m \int_0^t C_f(x,t') [k a_m \exp(-k a_m (t - t'))] dt' \quad (21)$$

and it can be verified by differentiation that Eq. (21) satisfies Eq. (15) when  $M(x,0) = 0$ , so that the solution for  $C_f(x,t)$  from Eqs. (1) and (21) is given by Eqs. (16) and (17).

To evaluate the mass transfer coefficient  $k$ , substitute for  $\bar{H}(t)$  the "long-time" infinite series solution<sup>5</sup> given by

$$\bar{H}(t) = 1 - \sum_{n=0}^{\infty} \frac{B}{(2n+1)^2 \pi^2} \exp \left[ \frac{-D_e (2n+1)^2 \pi^2 t}{4B^2} \right] \quad (22)$$

The resulting expression for  $M(x,t)$  is

$$M(x,t) = \phi_m R_m \int_0^t C_f(x,t') \left[ \frac{2D_e}{B^2} \sum_{n=0}^{\infty} \exp \left[ \frac{-D_e (2n+1)^2 \pi^2 t}{4B^2} \right] \right] dt' \quad (23)$$

Now, let  $k = 2\gamma D_e/B$ , where  $\gamma$  is a numerical constant, and note that  $a_m = 1/B$ . Comparing Eqs. (21) and (23), it can be seen that for an appropriate value of  $\gamma$ , on the order of 1 to  $\pi^2/6$ , Eq. (21) should provide a reasonable approximation for Eq. (23) when  $D_e t/B^2$  is large enough so that the series in Eq. (23) can be truncated after the first term. A value for the constant  $\gamma$  can be obtained as follows.

Let  $\bar{H}_b$  and  $\bar{H}_e$  denote the values of  $\bar{H}$  obtained from Eqs. (19) and (22), respectively; let

$$F_b = \frac{\frac{\partial \bar{H}_b}{\partial t}}{\frac{\partial \bar{H}_e}{\partial t}} ;$$

and again apply the mean value theorem to Eq. (6). The ratio of approximate to exact values of  $M(x,t)$  from Eqs. (21) and (23), respectively, is given by the term  $F_b^*$  defined analogously to  $F_a^*$  above. For very small  $D_e t/B^2$ , corresponding to small  $\bar{H}_e$ ,  $F_b \approx 0$ . As  $D_e t/B^2$  increases, and  $\bar{H}_e$  approaches unity,  $F_b$  will increase to some finite value which depends on the constant  $\gamma$ .

For a given value of  $\gamma$ , it appears reasonable to estimate  $F_b^*$  by the average  $\bar{F}_b$  of  $F_b$  with respect to  $\bar{H}_e$ , that is

$$\bar{F}_b = \frac{1}{\bar{H}_e} \int_0^{\bar{H}_e} F_b(\bar{H}) d\bar{H}$$

Values of  $\bar{F}_b$  as a function of  $\bar{H}_e$  are shown on Fig. 3 for  $\gamma = 1.0, 1.25, 1.50$ , and  $2.0$ . Since the values of  $\bar{F}_b$  corresponding to  $\gamma = 1.5$  show the least average relative deviation about 1.0, the value for  $\gamma$  was taken as 1.5, from which  $k = 2\gamma D_e/B = 3D_e/B$ .

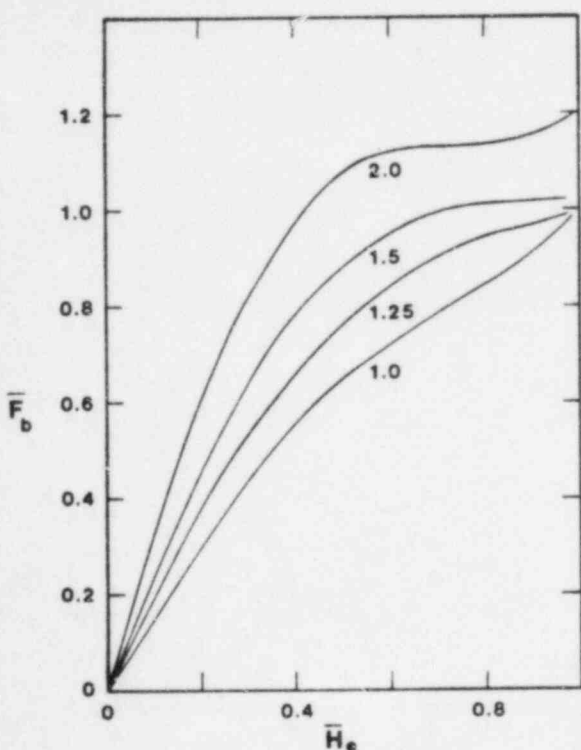


Fig. 3. Values of  $\bar{F}_b$  versus  $\bar{H}_e$  (numbers on curves are values for numerical constant  $\gamma$ ).

Criteria for applying the linear-driving-force approximation, that is calculating  $M(x,t)$  from Eq. (21) with  $k = 3D_e/B$ , will depend on the errors which are acceptable in  $M(x,t)$ . To define those

criteria, it is initially assumed, and later verified, that when the linear-driving-force approximation applies, the mean radionuclide residence time is given approximately by the maximum mean residence time  $\theta_m$  defined above. Again, let  $t^* = t - x/v$  and  $\theta_m^* = \theta_m - x/v$ . First, consider cases where  $t^* > \theta_m^*$ . From Fig. 3 ( $\gamma = 1.5$ ),  $\bar{F}_b \approx 1.0$  when  $\bar{H}_e \geq 0.7$  or  $D_e t^*/B^2 \geq 0.5$  (Table II). If radionuclide residence times do not vary greatly about  $\theta_m$ , and if  $\theta_m^* \geq 0.5B^2/D_e$ , perturbations in radionuclide concentrations would be exposed to a porous matrix throughout which  $\bar{F}_b \approx 1.0$ . For some cases, radionuclide residence times will vary significantly about the mean. If only small errors in radionuclide discharges are acceptable, a more conservative criterion is  $\theta_m^* \geq B^2/D_e$  or

$$\bar{x} = \frac{D_e m_f \phi_m R x}{v B^2} = \frac{\frac{D}{2} m_f \phi_m x}{v B^2} \geq 1 \quad (24a)$$

It seems reasonable to also define a less restrictive criterion since, as mentioned previously, parameter values in the expression for  $M(x,t)$  may often involve uncertainties of 20 to 30 percent. From Fig. 3,  $\bar{F}_b \geq 0.7$  when  $\bar{H}_e \geq 0.36$  or  $D_e t^*/B^2 \geq 0.1$  (Table II). Therefore, if  $\theta_m^* \geq 0.1B^2/D_e$ , perturbations in radionuclide concentrations would be exposed to a porous matrix throughout which  $0.7 \leq \bar{F}_b \leq 1.0$ . The relative errors in  $M(x,t)$  would be similar to or less than uncertainties in parameter values. Again, allowing for variations in radionuclide residence times about  $\theta_m$ , a reasonable criterion would be  $\theta_m^* \geq 0.2B^2/D_e$  or

$$\bar{x} = \frac{D_e m_f \phi_m R x}{v B^2} = \frac{\frac{D}{2} m_f \phi_m x}{v B^2} \geq 0.2 \quad (24b)$$

when errors of 20 to 30 percent in  $M(x,t)$  and the corresponding errors in  $C_f/C_0$  are acceptable.

Now consider cases where  $t^* < \theta_m^*$ . For a given ratio  $D_e \theta_m^*/B^2$ , the values of  $\bar{F}_b$  will decrease as  $D_e t^*/B^2$  becomes less than 0.5, while the relative errors in  $M(x,t)$  and the corresponding errors in  $C_f(x,t)/C_0$  will increase. If  $D_e \theta_m^*/B^2$  is large, those errors will occur when  $C_f/C_0 \approx 0$ , since for a reasonable variation in radionuclide residence times about  $\theta_m$ ,  $C_f/C_0$  will be negligible when  $t^* \ll \theta_m^*$ . If  $D_e \theta_m^*/B^2$  is of the order of 1 or less, relative errors in  $M(x,t)$  and  $C_f(x,t)/C_0$  will occur when values of  $C_f/C_0$  are significantly greater than zero. Therefore, when  $\theta_m^*$  is large, the criteria given by Eqs. (24a) and (24b) should be applicable for any value of  $t$ . When  $D_e \theta_m^*/B^2$  is of the order of 1 or less, the criteria are also applicable, but the errors in radionuclide discharges calculated using the approximation will increase substantially as  $D_e t^*/B^2$  becomes less than about 0.2.

**Equivalent Porous Medium.** In this approximation, it is assumed that relaxation times, denoted by  $t_r$ , for perturbations in radionuclide concentrations in the porous matrix are small relative to the time scale of interest. This assumption implies that the fracture fluid and porous matrix are in local equilibrium with respect to radionuclide diffusion. In which case,  $\bar{H}(t)$  in Eq. (6) approaches unity during a time interval which is less than the time required for



$C_f(x,t)$  to change appreciably, and  $M(x,t) = \phi_m R_m C_f(x,t)$ , which is obtained by integrating by parts in Eq. (6) and noting that  $C_f(x,0) = 0 = H(0)$  for  $x > 0$ . Eqs. (1) and (2) reduce to a single equation involving only  $C_f$

$$\frac{\partial C_f}{\partial t} + \frac{v}{R_f} \frac{\partial C_f}{\partial x} = 0 \quad (25)$$

which has the solution

$$C_f(x,t) = C_0 S\left(t - \frac{R_f x}{v}\right) \quad (25)$$

where  $S(t)$  denotes the Heaviside unit step function.

The validity of the equivalent-porous-medium approximation depends on the relaxation time  $t_r$  and the mean radionuclide residence time. Define  $t^*$  and  $\theta_m^*$  as above. Again, assume that the mean radionuclide residence time is approximately the maximum residence time  $\theta_m$ . If  $t_r \ll \theta_m^*$ , then perturbations in radionuclide concentrations would be exposed to fracture fluid and porous matrix near equilibrium. For purposes here, take  $t_r$  as the time required for  $H(t^*)$  to become approximately unity. From Table II,  $t_r = 2B^2/D_e$ , and a criterion for applying the equivalent-porous-medium approximation is  $\theta_m^* = (x/v)/(R_f - 1) \gg t_r = 2B^2/D_e$ . Allowing for perturbations in concentration and heterogeneity in the system, a reasonable criterion would be<sup>9</sup>

$$\bar{X} = \frac{D_e m_f \phi_m R_m x}{v B^2} = \frac{\frac{D}{2} m_f \phi_m x}{v B^2} \geq 50 \quad (27)$$

Provided that the above criterion is satisfied, the mean radionuclide residence time does correspond to the maximum residence time  $\theta_m$  since the porous matrix and fracture fluid essentially are near equilibrium. The actual errors in  $C_f(x,t)$  which result from applying the equivalent-porous-medium approximation are discussed below.

#### EVALUATION OF CRITERIA AND APPROXIMATIONS

To evaluate the above approximations, dimensionless breakthrough curves,  $C_f/C_0$  versus  $\bar{V}/\bar{X}$  were calculated using each of the three approximations. The approximate curves were then compared with the exact solution given by Eq. (3). The infinite integral in Eq. (3) was evaluated numerically using the method discussed by Rasmuson and Merethnieks.<sup>8</sup> Results for  $\bar{X} = 0.2, 1.0$ , and 50 are shown on Figs. 4 through 6, respectively. For convenience in discussing these results, note that  $(t - x/v)/(B^2/D_e) = \bar{V}/2$  and  $(\theta_m - x/v)/(B^2/D_e) = \bar{X}$ ; therefore  $\bar{V}/\bar{X} = 2(t - x/v)/(\theta_m - x/v) = 2t^*/\theta_m^*$ .

First consider the semi-infinite-medium approximation. Fig. 4 shows that for  $\bar{X} = 0.2$ , the approximation provides an excellent estimate for the exact solution when  $\bar{V}/\bar{X}$  is less than 3, and at larger values of  $\bar{V}/\bar{X}$  the relative errors in  $C_f/C_0$  due to the approximation are small. Therefore, the more restrictive criterion given by Eq. (14a) appears valid, and as previously predicted, the errors resulting from using the approximation are not significant until

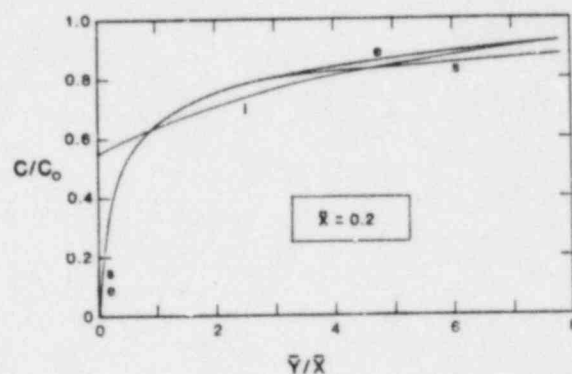


Fig. 4. Comparison of radionuclide discharges calculated from approximate and exact solutions when  $\bar{X} = 0.2$  (e, l, and s denote results for exact solution, linear-driving-force approximation, and semi-finite-medium approximation, respectively).

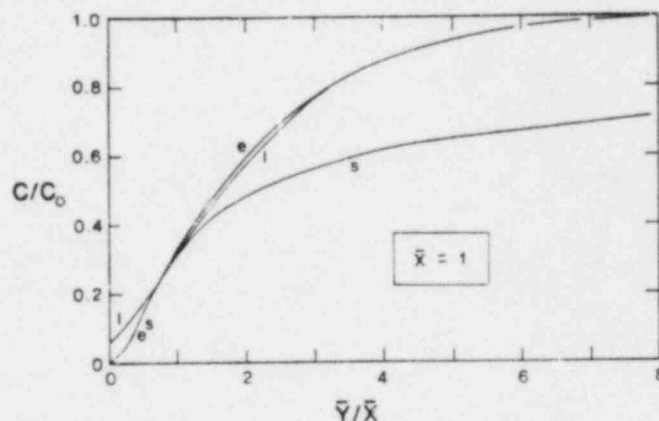


Fig. 5. Comparison of radionuclide discharges calculated from approximate and exact solutions when  $\bar{X} = 1.0$  (e, l, and s denote results for exact solution, linear-driving-force approximation, and semi-finite-medium approximation, respectively).

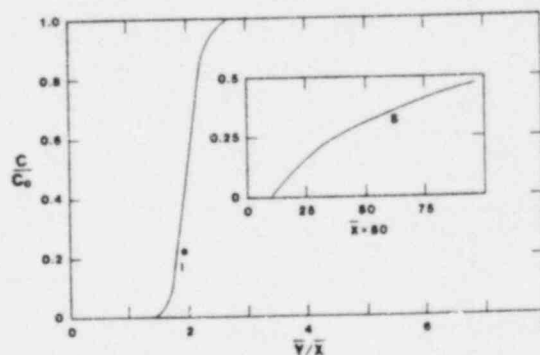


Fig. 6. Comparison of radionuclide discharges calculated from approximate and exact solutions when  $\bar{X} = 50$  (e, l, and s denote results for exact solution, linear-driving-force approximation, and semi-finite-medium approximation, respectively).

$t > \theta_m$  ( $t = \theta_m$  when  $\bar{Y}/\bar{X} = 2$ ). Fig. 5 shows that for  $\bar{X} = 1.0$ , the approximation provides an excellent estimate for the exact solution when  $\bar{Y}/\bar{X} < 1$ . As  $\bar{Y}/\bar{X}$  approaches 2 (or  $t/\theta_m = 1$ ), the relative error between approximate and exact solutions is about 20 to 30 percent, which is consistent with the errors previously predicted. Furthermore, the relative error does not increase appreciably as  $\bar{Y}/\bar{X}$  (or  $t/\theta_m$ ) becomes large. Therefore, the less restrictive criterion given by Eq. (14b) appears valid, provided that the resulting errors in  $M$  and  $C_f/C_0$  are acceptable. For  $\bar{Y}/\bar{X} > 1$ ,  $C_f/C_0$  is underestimated because  $M$  is overestimated. Fig. 6 ( $\bar{X} = 50$ ) shows that as  $\bar{X}$  becomes large, the semi-infinite-medium approximation becomes unacceptable.

Next, consider the linear driving force approximation. Figure 4 shows that for  $\bar{X} = 0.2$ , the approximation provides a reasonable estimate for the exact solution when  $\bar{Y}/\bar{X}$  is about 0.4 or greater, and the less restrictive criterion given by Eq. (15b) appears valid, provided that  $\bar{Y}/\bar{X} > 0.4$  (or  $t^*/(B^2/D_e) > 0.08$ ). Figure 5 shows that when  $\bar{X} = 1$ , the linear-driving-force approximation provides a very good estimate for the exact solution when  $\bar{Y}/\bar{X}$  is again about 0.4 or greater. At lesser values of  $\bar{Y}/\bar{X}$ , the relative error between approximate and exact solutions increases substantially, but the actual concentrations  $C_f/C_0$  from either solution are small. Therefore, the more restrictive criterion given by Eq. (15a) should be generally applicable provided that small concentrations at early times need not be estimated accurately. On Fig. 6 ( $\bar{X} = 50$ ), the curves for the linear-driving-force approximation and exact solution essentially coincide, which shows that the approximation provides an excellent estimate for the exact solution at large values of  $\bar{X}$ .

Also, from Figs. 5 and 6, it can be seen that at large  $\bar{X}$ , the mean radionuclide residence time is essentially  $\theta_m$ .<sup>1</sup> As  $\bar{X}$  becomes small, the deviation in residence times about the mean increases, but the mean residence time remains on the order of  $\theta_m$ .

Finally, consider the equivalent-porous-medium approximation. In Figs. 4-6, the approximate solution given by Eq. (26) corresponds to a vertical line at  $\bar{Y}/\bar{X} = 2$  (or  $t = \theta_m$ ). The errors associated with equivalent-porous-medium approximations have been discussed previously.<sup>9</sup> The essential features of those errors can be seen on Figs. 4-6. In particular, as breakthrough occurs ( $C_f/C_0$  becomes nonzero), the exact solution appears to be "dispersed" about the solution for the equivalent porous medium. As  $\bar{X}$  becomes large, that apparent dispersion becomes smaller, and at sufficiently large  $\bar{X}$  and  $\bar{Y}/\bar{X}$ , would have negligible effect on cumulative radionuclide discharges. For example, when  $\bar{X} = 50$ , the error in cumulative radionuclide discharges calculated using the equivalent-porous-medium approximation will be small provided that  $\bar{Y}/\bar{X}$  is about 2.6 or greater. Therefore, the criterion given by Eq. (27) appears valid provided that the time period of interest corresponds to  $\bar{Y}/\bar{X}$  less than about 1.4 or greater than about 2.6.

#### APPLICATIONS

Approximate methods for calculating radionuclide transport can be very useful in performance assessment studies. The approximations described above can easily be incorporated into transport codes and used to obtain realistic estimates of radionuclide releases.

The NWFT/DVM computer code<sup>10</sup> was developed at Sandia National Laboratories to simulate contaminant

transport for performance assessment studies. The program represents a known velocity field as a simplified network of one-dimensional transport segments. It can model the transport of radionuclide decay chains of any length, with isotopes having different retardation factors, and with various types of source terms. A new version of NWFT/DVM<sup>11</sup> treats flow and transport through fractured, porous media. Advection is assumed to take place in a set of parallel fractures and radionuclides diffuse into the adjoining rock matrix. Both the linear-driving-force and the equivalent-porous-media approximations are available in this version of the code. The analytical solutions derived in the previous sections were used to benchmark the linear-driving-force approximation of this computer code.

Dimensionless breakthrough curves were calculated with NWFT/DVM for several values of  $\bar{X}$  using the linear-driving-force approximation. Parameter values used in the calculations are listed in Table III. Two sets of calculations were carried out to determine the effect of flow velocity on the breakthrough curves generated. In order to simulate breakthrough for different values of  $\bar{X}$ , discharges were calculated for several values of path length  $x$ . Representative results are compared to the exact analytical solution Eq. (3) and the analytical solution for the linear-driving-force assumption Eq. (16). The numerical solution of NWFT/DVM agrees well with the analytical solutions for both the high and low velocity cases for  $\bar{X} = 1$  and  $\bar{X} = 50$  (Figs. 7-9). In addition, the numerical linear-driving-force approximation of NWFT/DVM agrees well with the exact analytical solution for  $\bar{X} = 0.2$  when  $\bar{Y}/\bar{X} > 2$ .

Table III

Parameter Values for NWFT/DVM Calculations

Fracture Aperture $2b$	100 $\mu\text{m}$
Fracture Spacing $2(B+b)$	Case 1: 10 cm
	Case 2: 50 cm
Fracture Porosity $\phi_f$	Case 1: $10^{-3}$
	Case 2: $2 \times 10^{-4}$
Matrix Porosity $\phi_m$	0.01
Tortuosity $\alpha^2$	10
Molecular Diffusion Coefficient $D$	$1.6 \times 10^{-5} \text{ cm}^2/\text{s}$
Matrix Retardation Factor $R_m$	1.0
Fluid Velocity $v$	Case 1: 10 cm/day
	Case 2: 0.75 cm/day

Eqs. (24a), (24b), and (27) were used to identify geochemical and hydrological conditions under which the semi-infinite-medium, linear-driving-force, and equivalent-porous-medium approximations are valid. Fig. 10 illustrates the application of these criteria to site-specific data for tuff<sup>12,13</sup> and basalt<sup>14,15</sup> and generic data for granite.<sup>16,17,18</sup> The plotted points bracket ranges of hydrological and geochemical parameters that are representative of these media. The parameter values used in constructing the plots have been tabulated by Erickson and others.<sup>19</sup> Lines representing  $\bar{X}$  values of 0.2, 1, and 50 divide the graph into regions within which at least one of the approximations will provide acceptable results. It can be seen that for tuff, the equivalent-porous-medium approximation should usually be valid even for relatively thin beds ( $x = 30 \text{ m}$ ). For basalt and granite, the semi-infinite-medium approximation or



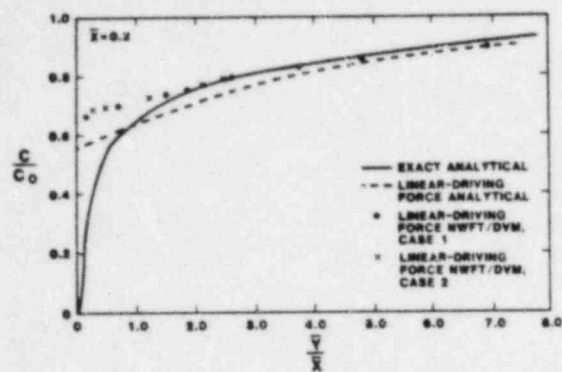


Fig. 7. Comparison of radionuclide discharges calculated with analytical exact solution and linear-driving-force approximation when  $\bar{X} = 0.2$ . For case 1, the fluid velocity = 10 cm/day and the distance from the source,  $x$ , is 36 cm. For case 2, fluid velocity = 0.75 cm/day and  $x = 14$  cm. (See Table III.)

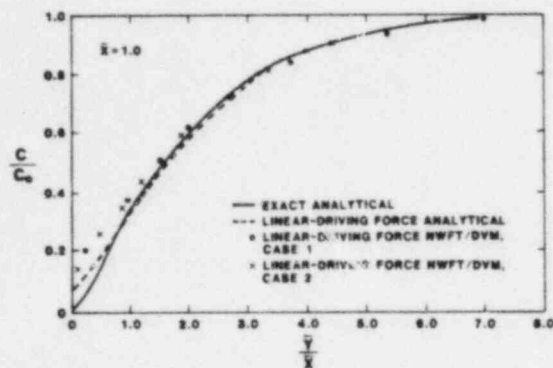


Fig. 8. Comparison of radionuclide discharges calculated with analytical exact solution and linear-driving-force approximation for  $\bar{X} = 1.0$ . For case 1,  $x = 180$  cm; for case 2,  $x = 68$  cm. See Table III for other parameter values.

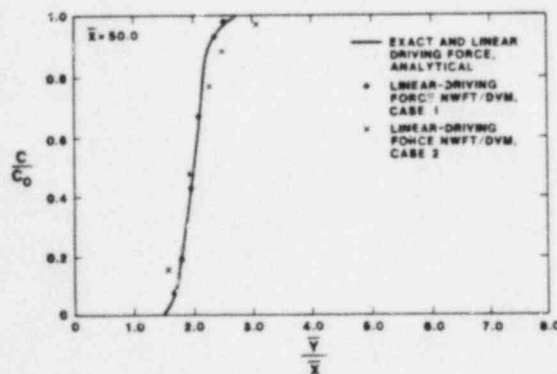


Fig. 9. Comparison of radionuclide discharges calculated with analytical exact solution and linear-driving-force approximation for  $\bar{X} = 50$ . For case 1,  $x = 90.5$  m; for case 2,  $x = 34$  m. See Table III for other parameter values.

the linear-driving-force approximation may be required for most calculations.

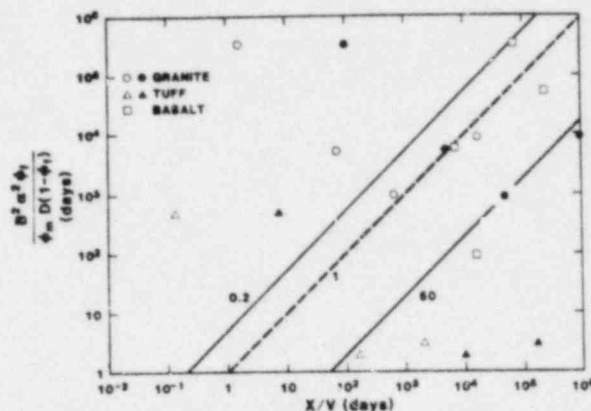


Fig. 10. Application of criteria to representative site-specific data for granite, basalt, and tuff. Numbers on lines are values of  $\bar{X}$ . Areas below lines marked '0.2' and '50' correspond to conditions under which linear-driving-force and porous-medium approximations, respectively, apply. The semi-infinite-medium approach applies in the area above the line marked '1'. Solid and open symbols refer to transport distances of  $x = 2000$  m and  $x = 30$  m, respectively.

#### CONCLUSIONS AND RECOMMENDATIONS

The results above are encouraging and indicate that the semi-infinite-medium, linear-driving-force, and equivalent-porous-medium approximations could be useful for performance assessment of HLW repositories in fractured, porous rock. The radionuclide discharges calculated by the linear-driving-force approximation used in the finite-difference code NWFT/DVM agree well with those calculated using an exact analytical solution for a range of hydrological parameters. Furthermore, the equivalent-porous-medium approximation could extend the results from Sandia's Geochemical Sensitivity Analysis Program to systems involving fractured, porous rock. However, additional evaluation is necessary. Cases must be examined in which the radionuclide material balances include terms for chemical reactions, radioactive decay, and production of nuclides. In addition, an assessment should be made of the sensitivity of the approximations to heterogeneities in fracture spacing, aperture and geometry and the presence of fracture-fill minerals.

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